

Assessment of glacier water resources based on the Glacier Inventory of China

KANG Ersi, LIU Chaohai, XIE Zichu, LI Xin, SHEN Yongping

*Cold and Arid Regions Environmental and Engineering Research Institute, Chinese Academy of Sciences, Lanzhou 730000, China
E-mail: eskang@lzb.ac.cn*

ABSTRACT. According to the division into subareas of water-resource distribution in China, and based on the Glacier Inventory of China (GIC), China's total glacier water storage is $5040.2 \times 10^9 \text{ m}^3$, 33.0% of which is distributed in the southwest drainage basins and 64.1% in the northwest inland drainage basins, forming enormous solid reservoirs with plentiful freshwater storage. Glacier change in China is estimated for the periods from the Little Ice Age (LIA) maximum to about 1960 and from 1960 to 1995. The relative glacierized area loss is 23% and 8.9% respectively for maritime glaciers and 15% and 4.9% respectively for continental glaciers. The normal annual glacier meltwater runoff is estimated at $60\,465 \times 10^6 \text{ m}^3$ by the climate parameter temperature-index method, 38.7% of which is distributed in the northwest inland drainage basins, and at $61\,574 \times 10^6 \text{ m}^3$ by the glacier system temperature-index method, 41.5% of which is distributed in the northwest inland drainage basins. Simulation of glacier meltwater runoff under the temperature change ratio of 0.03 K a^{-1} by the glacier system model in west China between 1980 and 2000 indicates that the total glacier meltwater runoff increment is 10.8% overall, 14.3% in the inland drainage basins of northwest China and 9.0% in the outflow drainage basins of southwest China.

1. INTRODUCTION

Glaciers are not only active in energy and mass exchange with the atmosphere, but are also important water resources that contribute meltwater to river discharge. Glaciers participate in the global water cycle and, with their solid water storage, are an important component of the water balance. As solid reservoirs, glaciers continue to receive the mass nourishment of solid precipitation from the atmosphere, and their meltwater feeds and regulates river discharge. Their solid water storage can become usable water resources only when it is melted and generates meltwater runoff (Kang and others, 2008). According to the published 12 volumes of the Glacier Inventory of China (GIC) consisting of 22 parts in 21 books, there are 46 377 glaciers covering an area of $59\,425 \text{ km}^2$, with ice storage of 5600 km^3 , in western China (Shi and others, 2008). Total glacier meltwater runoff in China is estimated at $60.465 \times 10^9 \text{ m}^3 \text{ a}^{-1}$ (Yang, 1991). Thus, in the integrated assessment of water resources in China, the assessment of glacier water resources is very important.

To support the rational allocation and management of water resources, an assessment of China's glacier water resources must be made for the water-resource subareas. First, the glaciers and their water-resource status must be summed up in different orders in west China, the assessment including the distribution, types, characteristics and water-balance status of the glaciers, and their role in the mountain runoff of glacial meltwater. Secondly, a trend analysis of glacier change must be carried out to show historical variations, and the future trend of glaciers and their meltwater runoff predicted under various climate-change scenarios, especially for the inland river basins where the glacial meltwater constitutes a large proportion of the mountain runoff. Thirdly, scientific understanding of interactions among glaciers, climate change and water resources must be improved.

The basis for assessing glacier water resources is the completion of the GIC (Shi and others, 2008). Statistical analyses are performed on the glacier inventory data to show the characteristics of glacier distribution. In addition, the results of the long-term glacier monitoring at Tien Shan Glaciological Station are analyzed, and data from field expeditions and investigations on representative glaciers are used. Then the China Glacier Information System (Wu and Li, 2004) is applied and some simulation methods are used. Based on the above, we have made a preliminary assessment of China's glacier water resources.

This paper presents the main results of the assessment based on the GIC, and discusses the distribution of glaciers in the water-resource subareas, glacier change, normal glacier runoff and changes in glacier runoff.

2. DISTRIBUTION OF GLACIERS IN THE WATER-RESOURCE SUBAREAS

According to the manual of the World Glacier Inventory, China's mountain glaciers are distributed in the drainage basins which are classified into five orders and ten first-order drainage basins (Shi and others, 2008). According to the Chinese Hydraulic Engineering Society (2002), glaciers in west China are distributed in four first-order water-resource subareas: the Northwestern Rivers region, the Southwestern Rivers region, the Yellow River region and the Yangtze River region (Fig. 1). The statistics of glacier distribution in the water-resource subareas are based on the GIC data. The results are shown in Figure 2.

Figure 2 shows that the glaciers are mainly located in the Northwestern Rivers region, which is an arid area consisting of various inland river basins (Liu and others, 1999), with most of the others located in the Southwestern Rivers region, which is a humid area with rivers draining to the ocean. The glaciers in the arid inland area are of continental type, while

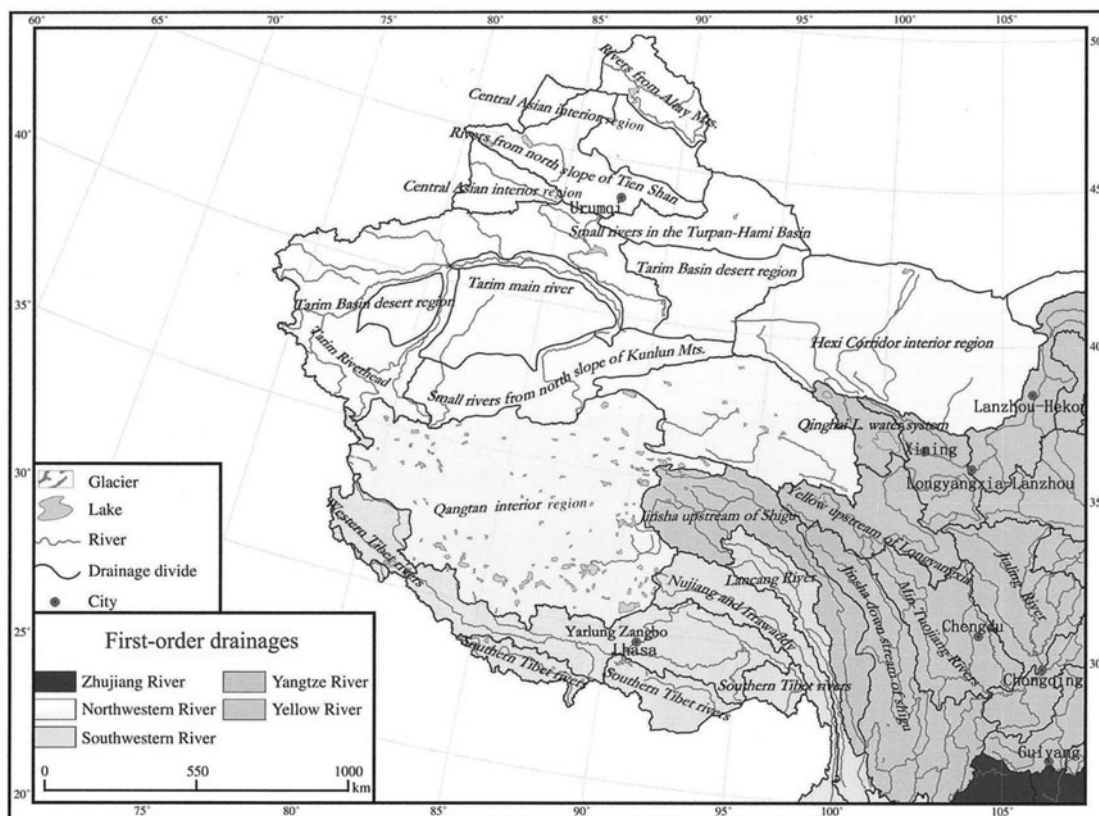


Fig. 1. First-order water-resource subareas of China (from CHES, 2002; Shi and others, 2008, p. 43).

those in the humid outflow area are of maritime temperate type (Shi and others, 2008, p.31–37). According to ice storage from the GIC, the total glacier water storage is $5040.2 \times 10^9 \text{ m}^3$, 64.1% of which is stored in the northwest inland drainage basins and 33.0% in the southwest drainage basins. Figure 2 also shows that the average scale of glaciers is slightly larger in the arid inland area than in the humid outflow area.

The number of glaciers in the arid northwest is 27 024, covering an area of $35\,468.56 \text{ km}^2$, with an ice storage of 3574.75 km^3 . The area is divided into ten second-order water-resource subareas. Table 1 shows the distribution of glaciers in the subareas. 50% of the glacierized part of the area is distributed in the mountains surrounding the Tarim basin, which covers 45% of the arid area, and 22.1% is distributed in the northwest Qinghai–Tibetan Plateau (Qiangtang plateau), which covers 31% of the arid area (Shi and others, 2008). The other subareas, covering 24% of the arid area, contribute 27.9% of the glacierized area.

3. GLACIER CHANGE

3.1. Methodology

Completion of the GIC took more than 20 years, and the time-span of the photogrammetric topographic maps used for the inventory is large, from 1956 to 1984. To investigate glacier change, the Little Ice Age (LIA) moraine distribution was investigated with air photos, and comparative studies were carried out by field measurement on representative glaciers and topographic maps from different times. Repeated measurements of glacier variations were also implemented by remote sensing, stereophotogrammetric

survey and aerial survey. The glacierized area of the LIA maximum can be estimated from the moraine sediment in the representative regions (S. Liu and others, 2002). For the present glacier change, there are various short-term field investigations and measurements in representative glacierized areas, and perennial glaciological measurements have been carried out at the source area of the Ürümqi river since the end of the 1950s (S. Liu and others, 2002).

According to statistical analysis of the existing glaciological measurements both at the LIA moraine and at existing glaciers, the percentage change in glacierized area is strongly correlated with the scale of the glaciers, and is also dependent on the scale class of the glacier area (Liu and others, 1999). The relationships are as follows (S. Liu and

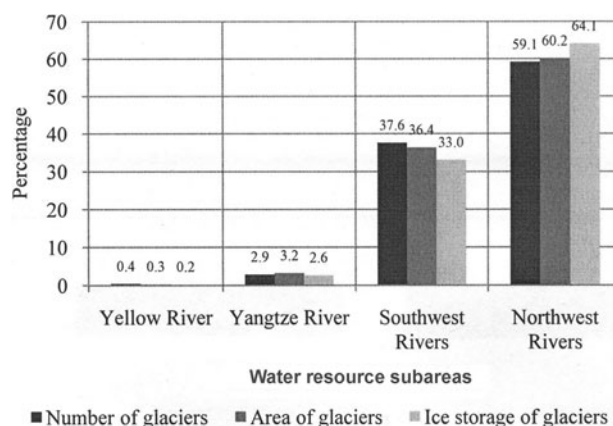


Fig. 2. Distribution of glaciers in the first-order water-resource subareas of China (from Shi and others, 2008, p. 44).

Table 1. Glacier distribution in the second-order water-resource subareas in the northwest arid inland area of China (data from GIC)

Water-resource subarea	Number of glaciers		Area		Ice storage	
		%	km ²	%	km ³	%
Hexi inland basins	2194	8.1	1334.77	3.8	61.54	1.7
Qinghai lake	22	0.1	13.29	0.0	0.59	0.0
Qaidam basin	1559	5.8	1851.76	5.2	128.07	3.6
Turpan-Hami basin	446	1.7	252.73	0.7	12.63	0.4
Southern Altai	13	0.0	3.91	0.0	0.10	0.0
Central and West Asia inland basins	2385	8.8	2048.16	5.8	143.71	4.0
Northern Tien Shan	3399	12.6	2250.19	6.3	137.33	3.8
Tarim basin	10297	38.1	17745.51	50.0	2139.94	59.9
North flank of Kunlun	1368	5.1	2132.14	6.0	173.34	4.8
Qiangtang plateau	5341	19.8	7836.10	22.1	777.48	21.7
Total	27 024	100.0	35 468.56	100.0	3574.75	100.0

others, 2002, p. 18–20):

$$S_L = \frac{S}{1 - \Delta S/S_L}, \quad (1)$$

where S_L is the area of a glacier during the LIA maximum (km²), S is the area of an existing glacier (km²) and ΔS is the area reduction of a glacier between the LIA maximum and 1960 (km²). The glacier area reduction $\Delta S/S_L$ has a good correlation with the scale of the glaciers as:

$$\frac{\Delta S}{S_L} = aS_i^{-b}, \quad (2)$$

where a and b are regional positive constants and S_i is the mid-value of an area-scale series of glaciers in a basin or a region (km²). Table 2 shows the correlation parameters of Equation (2) in the representative glacierized basins.

Thus the glacier change is estimated with the statistical relationships based on the GIC data in the inland arid area of northwest China and in the outflow area of southwest China from the LIA maximum to 1960, when the status of existing glaciers can be obtained from the GIC. From the field measurements and investigations of glacier change in the representative glacierized areas from 1960 to 2000, it is found that the relationship of area reduction to the scale of the glaciers can also be simulated by Equations (1) and (2) (S. Liu and others, 2002, p. 21–33). Glacier change in China during recent decades is then estimated.

3.2. Glacier change since the LIA maximum and in recent years

At the LIA maximum, the glacier-covered area in west China was 75 438.4 km², which is 16 013.2 km² more than the GIC glacierized area, the percentage loss being 21.2% (Shi and others, 2008). The continental-type glaciers in the arid northwest contribute 20.7% of the loss, while the glaciers of maritime temperate type in the outflow area contribute 22.0%.

It is estimated that glaciers in west China underwent a 5.5% area loss during the period 1960–2000 (Table 3). China's glaciers have been divided into three types: maritime temperate, subcontinental and extreme continental (Shi and others, 2008, p. 31–33). The largest relative area loss by type is 8.9% for maritime temperate glaciers, next is 6.0% for subcontinental glaciers, and the smallest loss is 2.5% for extreme continental glaciers. The average area loss for continental-type glaciers is 4.5%.

The change in China's glacierized area since the LIA maximum shows large spatial variation. However, this change is characterized by the regularity that the percentage change is larger in the east and west border regions of China's glacierized area, closer to the vapour current source direction: the east closer to the east Asia monsoon and the west closer to the westerly vapour current from the Atlantic Ocean. The glacierized area loss is 37% in the Shiyang river basin, east Qilian Shan, and 35% in the Ertix

Table 2. Correlation of the percentage area loss of glaciers ($\Delta S/S_L$) to their scale (S_i) in the representative glacierized river basins from the LIA maximum to 1960 (R is correlation coefficient) (S. Liu and others, 2002, p. 19)

River basin	S_i						Correlation
	≤0.5 km ²	0.5–2 km ²	2–5 km ²	5–10 km ²	10–100 km ²	>100 km ²	
Shiyang river of east Qilian Shan	44.50	35.10	27.80				$\Delta S/S_L = 35.30S_i^{-0.18}$ ($R = 0.99$)
Liugouquan river of middle Qilian Shan	25.00	16.90	13.60				$\Delta S/S_L = 17.02S_i^{-0.27}$ ($R = 0.99$)
Hotan river of Kunlun Shan	12.30	10.60	9.00	6.10	1.90	0.60	$\Delta S/S_L = 11.18S_i^{-0.51}$ ($R = 0.95$)
Yarkant river of Karakoram mountains				7.20	4.00	3.00	$\Delta S/S_L = 16.59S_i^{-0.28}$ ($R = 0.99$)
Toxkan river of south Tien Shan	25.20	18.00	13.80	9.10	6.80		$\Delta S/S_L = 19.60S_i^{-0.42}$ ($R = 0.98$)
Ürümqi river of middle Tien Shan	45.70	28.10	25.10				$\Delta S/S_L = 26.38S_i^{-0.29}$ ($R = 0.94$)
Burqin river of Altai mountains	37.10	21.20	9.60	3.80			$\Delta S/S_L = 18.30S_i^{-0.65}$ ($R = 0.96$)

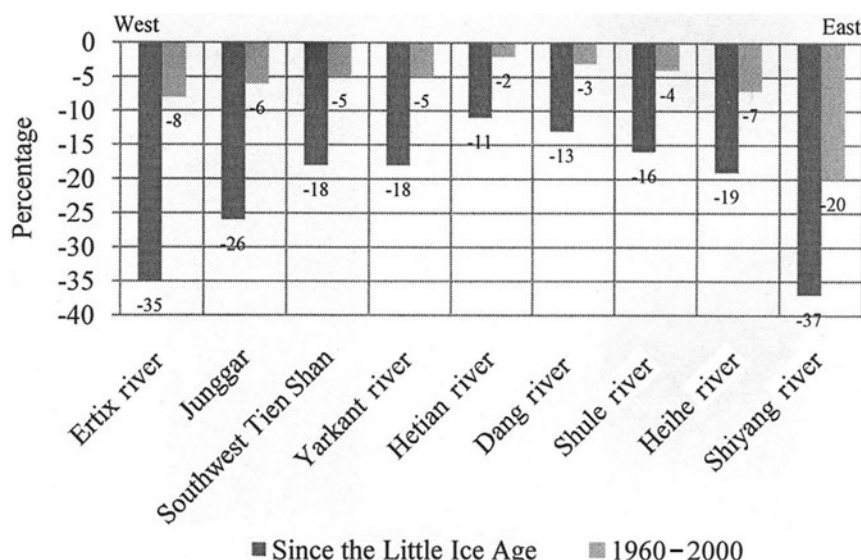


Fig. 3. Change of glacierized areas in northwest China from west to east since the LIA maximum and during recent years (from Shi and others, 2008, p. 176–180).

river basin, west Altai mountains (Fig. 3). From the east and west towards the central regions, which are further away from the vapour current source, precipitation decreases along with the percentage glacierized area, which reaches a minimum of about 10% in the west of the Qilian Shan and in the east of the west Kunlun Shan (Fig. 3). Statistical analysis shows that for glaciers with areas less than 0.5 km² the glacier area during the LIA maximum was twice that of existing glaciers in the GIC, while for glaciers with areas more than 10 km² it was only 1.03 times as much. Therefore, glaciers that have vanished since the LIA should be predominately small glaciers, while the large glaciers at the LIA maximum have experienced some area reduction but should still exist.

During the 40 year period 1960–2000, the area change of glaciers shows the same characteristics as the change since the LIA (Fig. 3). The percentage area loss increases spatially with increasing humidity, i.e. from the central region to the east and west boundary regions, which are closer to the water-vapour source direction (C. Liu and others, 2002; S. Liu and others, 2002). The highest percentage area loss occurs in the Shiyang river basin, east Qilian Shan, and the next highest losses occur in the Ertix river basin, west Altai mountains, and the Yili river basin, western Tien Shan (Fig. 3). The percentage area loss decreases with the increase in area and scale of an individual glacier. In the Tarim basin, which is surrounded by the Tien Shan, Pamir, Karakoram and Kunlun mountains, glaciers are larger, so the percentage area change is relatively small, less than the average for the total number of glaciers (Table 3).

A comparison between the glacier changes since the LIA maximum and those during recent years shows that the trend of glacier change since the LIA is dominated by glacial reduction, while in terms of the timescale, the percentage area loss greatly increases in recent years, reflecting global warming. The spatial variation of the relative area loss indicates that, along with the increase of continentality, from the east and west regions to the central regions of northwest China, the glaciers have become more stable and their percentage area loss has decreased.

4. ESTIMATION OF GLACIER RUNOFF

4.1. Climate parameter temperature-index method

Based on comparative measurements and experiments on glacial runoff at selected representative glaciers, glacial meltwater runoff in China is estimated by the temperature-index method, using the data published in the GIC (Yang, 1991, p. 72–73). Field measurements at more than ten of the fixed and semi-fixed stations at the representative glaciers in west China have shown that glacier melt is related to a power function of air temperature (Yang, 1981, 1982; 1991, p. 63–68; Shi and others, 2008, p. 160). The following equation is obtained:

$$\bar{a} = \varphi(T + 4.0)^{2.7}, \quad (3)$$

where \bar{a} is the mean daily water depth of glacial surface melt during the ablation period (mm d^{-1}), φ is a regional climate parameter, $\varphi = 0.382b$, where b is the ratio of net radiation to the incoming energy of energy balance (%), and T is the mean daily air temperature during the ablation period ($^{\circ}\text{C}$). The mean meltwater depth of glaciers is calculated at the median contour line, which divides the glacier surface into two parts of equal area. φ is obtained by the research on energy balance of glaciers in China, and T is calculated with an air-temperature–altitude gradient, taking the nearby standard meteorological stations as the base. In this way, a

Table 3. Glacier area change by type in west China from 1960 to 2000

Glacier type	Area		Area loss	
	1960 km ²	2000 km ²	km ²	%
Extreme continental	19 162.84	18 685.54	477.30	2.5
Subcontinental	27 008.18	25 390.53	1617.65	6.0
Maritime	13 254.16	12 076.28	1177.88	8.9
Total	59 425.18	56 152.35	3247.72	5.5

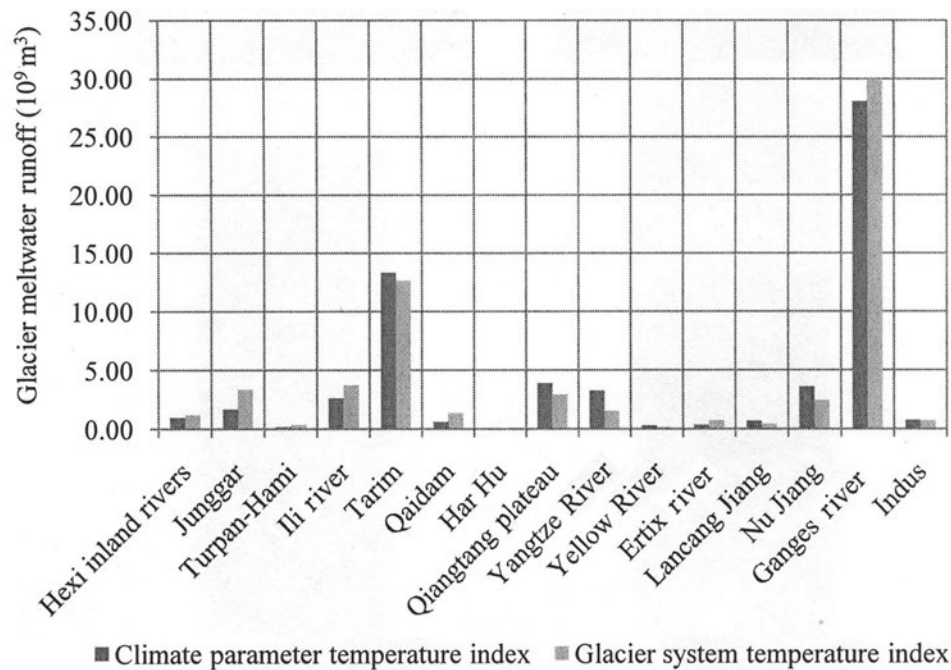


Fig. 4. Distribution of the normal annual glacier meltwater runoff (10^9 m^3) in China estimated by the climate parameter temperature-index method and the glacier system temperature-index method. The abscissa from Qiangtang plateau to the left indicates the inland drainage basins in the arid northwest, and that from the Yangtze River to the right indicates the outflow drainage basins (data from Shi and others, 2008, p. 1165, and from GIC).

preliminary estimate has been made of the annual mean volume of glacial meltwater runoff and its distribution in China (Yang, 1991; Kang and others, 2008).

4.2. Glacier system temperature-index method

Another temperature-index method is based on the glacier system concept (Kotlyakov and Smolyarova, 1990) and the summer season (June–August) temperature-index method (Kotlyakov and Krenke, 1982) expressed by Equation (4). Considering the glaciers in a region as a system, the glacier of median size in the area series of glaciers is taken to calculate the total melt of the glacier system in a region (Xie and Feng, 1996). Glacier annual total melt a (mm) can be calculated by the summer mean temperature t_s at the melt position on the glacier surface as

$$a = 1.33(9.66 + t_s)^{2.85}, \quad (4)$$

where t_s is calculated by

$$t_s = t_{pa} + \Delta h\gamma + \Delta t_j, \quad (5)$$

where t_{pa} is the temperature at the 600–500 mbar isobaric surface, Δh is the height difference between the isobaric surface and the glacier equilibrium line, γ is the temperature vertical lapse rate ($0.6\text{--}0.75 \text{ K}(100 \text{ m})^{-1}$) and Δt_j is the temperature jump value from the ground surface to the glacier surface, taken as -0.5 , 1.0 and 1.5 K according to the glacier area (Xie and others, 2006a). Then, taking the annual melt at the equilibrium line to represent the glacier total melt (Ohmura and others, 1992; Kang and others, 1994), the total regional glacier meltwater runoff can be obtained from the area distribution of the GIC.

4.3. Estimation of glacier meltwater runoff

From the above, the normal glacier meltwater runoff can be estimated with the normal air temperature. Figure 4 shows

the distribution of the normal glacier meltwater runoff in China estimated by the climate parameter temperature-index method (Yang 1991; Kang and others, 2008) and the glacier system temperature-index method (Xie and others, 2006a,b). The normal annual glacier meltwater runoff in China is estimated at $60.465 \times 10^9 \text{ m}^3$ by the climate parameter temperature-index method, 38.7% of which is distributed in the northwest inland drainage basins, and at $61.574 \times 10^9 \text{ m}^3$ by the glacier system temperature-index method, 41.5% of which is distributed in the northwest inland drainage basins. Thus, the glacier runoff results estimated by the two temperature-index methods are very similar. This suggests that both the regional climate parameter φ of Equation (3) and the glacier of median size of a glacier system for the regional glacier meltwater estimation by Equation (4) could well represent the regional characteristics, and that both Equations (3) and (4) could be applied to estimate the regional normal glacier meltwater runoff in west China.

In the inland drainage basins, glacier runoff is mostly distributed in the Tarim basin, and accounts for $13.342 \times 10^9 \text{ m}^3$, recharging 38.5% of the mountain runoff of the basin. In the outflow drainage basins, glacier runoff is mostly distributed in the Ganges river basin, and accounts for $28.048 \times 10^9 \text{ m}^3$, recharging 9.1% of the mountain runoff (Shi and others, 2008, p. 165).

5. CHANGE OF GLACIER RUNOFF

When temperature increases and precipitation remains unchanged, glacier runoff will increase and glacier area will reduce. When the glacier melt increase, caused by temperature rise, is more than the meltwater reduction caused by the decrease in glacierized area due to glacier retreat, glacier runoff will increase to a maximum value.

Based on this, Xie and others (2006a,b) developed a formulation to simulate the change of glacier runoff:

$$W = (\gamma_0 + \gamma_d)(S_0 - S_d), \quad (6)$$

where W is the annual glacier runoff volume, γ_0 is the glacier runoff depth in the start year on the glacier area S_0 , and γ_d is the increment of glacier runoff depth in the year when the glacier surface area is reduced by S_d . When glacier runoff increases first, then reduces to its value at the start year, S_d can be calculated by

$$S_d = \frac{S_0 \gamma_d}{\gamma_0 + \gamma_d}. \quad (7)$$

Suppose $\alpha = \gamma_d/\gamma_0$, then

$$S_d = \frac{S_0 \alpha}{\alpha + 1}. \quad (8)$$

According to the measurement on representative glaciers, there is a relationship between the average depth H (m) and area S (km²) of a glacier (Liu and Ding, 1986; Shi and others, 2008):

$$H = 53.21S^{0.3} - 11.32. \quad (9)$$

Thus, from Equations (6–9), a glacier system model is developed to estimate the change in glaciers and their runoff under temperature change (Xie and Feng, 1996; Xie and others, 2002, 2006a,b). Based on the GIC data, the change in glacier meltwater runoff can be estimated by the temperature-index method and the various temperature scenarios of climate change.

The simulation of glacier meltwater runoff under the temperature change of 0.031 K a⁻¹ by glacier system model in west China between 1980 and 2000 indicates that the total glacier meltwater runoff increment is 10.8%, and it is 14.3% in the inland drainage basins in northwest China and 9.0% in the outland drainage basins in southwest China.

6. CONCLUSION

Glaciers are an important water resource in west China, especially in the inland river basins in the arid northwest, where glacier meltwater is an important source of runoff recharge of mountain rivers. Based on the GIC, glacier water resources were assessed for glacier distribution in China's water-resource subareas, change of the glaciers, glacier runoff and its simulation under temperature change.

The glacier inventory only reflects the glacial status at the time when the glaciers were mapped. Application of the GIC data needs to take account of the changing status of the glaciers. In the assessment of China's glacier water resources, field measurements and investigations of representative glaciers are used to analyze the change of the glaciers to find some statistical relationships, then to estimate the glacier change from the GIC.

The glaciers recorded in the GIC represent the basic conditions of existing glaciers in China. The estimation of glacier meltwater runoff is based on the existing glaciers and the average temperature of time series for many years. Therefore, the estimated glacier runoff is the normal value. If glaciers are in stable and equilibrium conditions, the estimated glacier melt is the change in ice storage. Glacial retreat under rising temperature will increase ice melt, reducing the ice storage of the glaciers.

Estimation of glacier change and meltwater runoff based on the GIC is very preliminary, because very few glaciers

have been measured. Further field measurements should be made and research on glacier changes in response to climate change should be strengthened. The ongoing project of the second GIC with remote-sensing techniques will bring new progress in the assessment of glacier water resources and their changes in China.

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