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# The *Tropidoleptus carinatus* controversy: Did this brachiopod occur in the Devonian of the Paraná Basin, Brazil?

Roberto Videira-Santos<sup>1,2</sup>\* • and Sandro Marcelo Scheffler<sup>1</sup> •

<sup>1</sup>Universidade Federal do Rio de Janeiro, Museu Nacional, Departamento de Geologia e Paleontologia, Laboratório de Paleoinvertebrados – LAPIN, Quinta da Boa Vista, s/n, São Cristóvão, 20940-040, Rio de Janeiro, RJ, Brazil. <robvidsan@ufrj.br>, <schefflersm@mn.ufrj.br>

<sup>2</sup>Universidade Federal do Rio de Janeiro, Instituto de Geociências,Programa de Pós-Graduação em Geologia, Avenida Athos da Silveira Ramos, 274, Cidade Universitária, 21941-916, Rio de Janeiro, RJ, Brazil.

Non-technical Summary.—*Tropidoleptus* was a genus of brachiopods that had a very peculiar morphology and occurred in the Devonian period (about 419–358 million years ago) throughout the world. Although there have been reports of the presence of this genus in the Paraná Basin (Brazil) since the nineteenth century, its actual occurrence has been the subject of much debate. The objective of this work was to clarify this situation. For that, dozens of specimens from the Brazilian state of Mato Grosso were analyzed. The studied samples have typical characteristics of *Tropidoleptus*, therefore this genus actually also occurred in the Paraná Basin. *Tropidoleptus* possibly arrived in the Paraná Basin 385 million years ago, at the moment when there was a marine connection among the Brazilian Paraná, Parnaíba, and Amazonas basins. The occurrence of *Tropidoleptus* can be used as a marker of shallower marine environments, and in the Paraná Basin in particular, it can indicate that the rocks, where they are present, are less than 390 million years old.

**Abstract.**—*Tropidoleptus* is a brachiopod genus with a very peculiar morphology that existed throughout the Devonian as a cosmopolitan taxon. Although there have been reports of the presence of this genus in the Paraná Basin since the nineteenth century, its actual occurrence has been the subject of much debate. The objective of this paper was to clarify this situation. For this purpose, dozens of specimens from Monjolo dos Padres and Juscimeira outcrops (São Domingos Formation, possibly Givetian) in the Brazilian state of Mato Grosso were analyzed. The samples studied here exhibit the typical characteristics of *Tropidoleptus*, confirming that this genus also occurred in the Paraná Basin. It is possible that *Tropidoleptus* entered the Paraná Basin during the Eifelian/Givetian from the Amazonas and/or Parnaíba basins (Brazil), during a period of warm-water transgression. When *Tropidoleptus* is found in situ, it can serve as an indicator of a shallower water environment (shoreface to shallower portions of transitional offshore). In the Paraná Basin, its presence can indicate rocks no older than the Eifelian–Givetian transition.

### Introduction

The brachiopod *Tropidoleptus* underwent several morphological changes throughout its existence in the Devonian (Isaacson and Perry, 1977) and is considered a cosmopolitan genus (Harper, 2007). This genus has unique characteristics that make its classification into a distinct order challenging. It previously has been assigned to the orders Strophomenida, Terebratulida, and Orthida (see Hall and Clarke, 1894; Schuchert, 1929; Schuchert and Cooper, 1932; Williams and Wright, 1961; Wright, 1965). Jaanusson (1971) included *Tropidoleptus* in the order Telotremata, which would encompass spiriferids and terebratulids. On the other hand, Harper et al. (2003) and Harper (2007) suggested that this genus belong to an undefined order. More recently, Harper et al. (2010) proposed elevating the family Tropidoleptidae to the order Tropidoleptida.

In Brazil, the presence of this taxon in the Amazonas and Parnaíba basins is indisputably recorded (e.g., Rathbun, 1874, 1878; Carvalho, 1972; Fonseca and Melo, 1987; Melo, 1988;

However, its occurrence in the Paraná Basin is still a matter of debate (see Ribeiro, 2020). Although Derby (1895) documented the occurrence of *Tropidoleptus* in the Paraná Basin (Chapada dos Guimarães, Mato Grosso), he relied on a single poorly preserved specimen. Subsequently, Orville Derby himself issued a note in Clarke's (1913) monograph correcting his identification and stating that it was actually Leptocoelia sp. (= Australocoelia sp.). The specimen studied by Derby (1895) was not illustrated in his work, but Kunzler (2012) found it in the paleoinvertebrate collection of the Museu Nacional/UFRJ (MN 3323-I) and provided the first illustration of it. Upon review, Kunzler (2012) concluded that it was indeed Tropidoleptus, but due to the poor state of preservation of the specimen, its identification remains open to discussion (e.g., Ribeiro, 2020). Clarke (1895) reported the occurrence of Tropidoleptus? in samples from Jaguariaíva, Paraná state.

Gama Junior, 2008; Queiroz et al., 2013; Santos et al., 2021).

Caster (1947) reported the presence of specimens of *Tropidoleptus* in the Paraná Basin, specifically in the state of Mato Grosso, but neither described nor illustrated them. Almeida (1954) mentioned the occurrence of *Tropidoleptus* in the region of Chapada dos Guimarães. Suárez-Riglos (1967) mentioned



<sup>\*</sup>Corresponding author.

that Caster's material included specimens of Tropidoleptus. Additionally, Lange and Petri (1967) noted that Tropidoleptus carinatus (Conrad, 1839) was the most common macrofossil of the São Domingos "member", clearly referring to the Middle Devonian layers of Mato Grosso, as evident from the subsequent fossil listing. Finally, Quadros and Melo (1989) claimed to have found *Tropidoleptus* in the locality of Monjolo (dos Padres?) (Chapada dos Guimarães), but they did not provide descriptions or illustrations of the specimens.

As can be seen, the occurrence of Tropidoleptus in the Paraná Basin has been supported primarily by citations, with only one poorly preserved specimen illustrated and partially described (Kunzler, 2012). Therefore, the purpose of this paper is to definitively establish the presence of Tropidoleptus in the Paraná Basin by analyzing numerous samples from the state of Mato Grosso. Resolving this question is crucial because Tropidoleptus is a significant taxon for paleobiogeographic discussions concerning the connections between the main Brazilian intracratonic basins (Amazonas, Parnaíba, and Paraná) and other basins in South America, the Old World, and Laurasia during the Devonian (see Isaacson and Perry, 1977; Harper et al., 2010).

### **Geological setting**

The Paraná Basin, covering an area of about 1,500,000 km<sup>2</sup>, is among the largest intracratonic basins in South America. It extends across parts of Brazil, Argentina, Paraguay, and Uruguay (Melo, 1988; Grahn, 1992; Milani et al., 2007; Fig. 1). Devonian outcrops are found along the borders of this basin, specifically in the Brazilian states of Paraná (eastern border), Mato Grosso, Mato Grosso do Sul, and Goiás (northwestern border) (Melo, 1988). Although there are different proposals to divide the Devonian in the Paraná Basin into various different units (e.g., Bergamaschi, 1999; Milani et al., 2007; Grahn et al., 2010; Sedorko et al., 2018), this paper adopts the lithostratigraphic framework presented by Grahn et al. (2013) for the entire basin. In their classification, Grahn et al. (2013) divided the Devonian strata in the Paraná Basin into the Furnas Formation (Lochkovian), Ponta Grossa Formation (late Pragian-early Emsian), and São Domingos Formation (late Emsian-Frasnian).

São Domingos Formation.—The specimens examined in this formation were collected from the Monjolo dos Padres and Juscimeira outcrops, both located in the state of Mato Grosso (Fig. 1). Due to the friable sandstone lithology of these outcrops, dating them through palynology is not possible. However, based on the macrofaunal content, it is likely that these specimens are of Givetian age (see Quadros and Melo, 1989; Sedorko et al., 2021), suggesting their affiliation with the São Domingos Formation (Figs. 1–3). The São Domingos Formation is stratigraphically situated above the Ponta Grossa Formation and below the Itararé Group (Carboniferous) on the eastern border, and below the Aquidauana Formation (Carboniferous) on the northwest border. It represents a marine deposit that encompasses foreshore to offshore environments, with deltaic facies observed at some localities (Sedorko et al., 2018, Ribeiro et al., 2019). Lithologically, the São Domingos Formation is composed of poorly sorted conglomeratic sandstones, sporadically bituminous shales, and micaceous siltstones containing abundant plant fragments (Grahn et al., 2013). In the lower part of the São Domingos Formation (late Emsian-Eifelian), elements of the depleted Malvinokaffric fauna (or Malvinoxhosan, sensu Penn-Clarke and Harper, 2021) are present, while after the Givetian, organisms of the typical Malvinokaffric fauna become scarce or disappear (Bosetti et al., 2010, 2012).

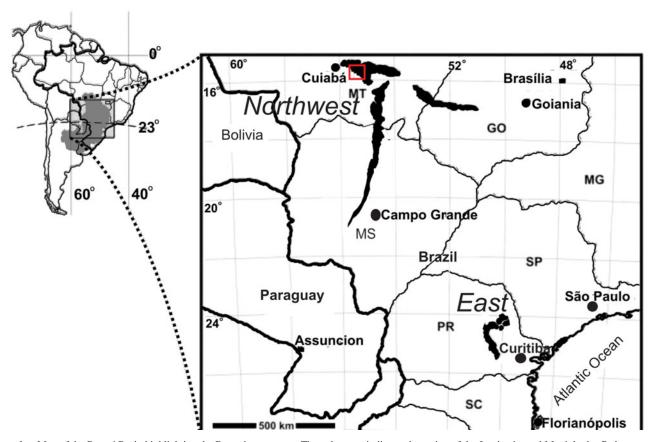
#### Materials and methods

The specimens examined in this study are housed in the paleoinvertebrate collection of the Museu Nacional/Universidade Federal do Rio de Janeiro (MN-I) and the Universidade Federal de Mato Grosso (MP, CD). The Museu Nacional specimens are the same ones previously analyzed by Caster (1947) and are part of the Caster Collection. Initially, these specimens were sent to the University of Cincinnati and later repatriated to the Museu Nacional collection in 2016 (see Scheffler et al., 2021). The specimens that were not affected by the tragic fire in 2018 because they were not inside the palace are MN 12002-I, MN 12003-I, MN 12004-I, MN 12005-I, MN 12006-I, MN 12007-I, MN 12008-I, MN 12009-I, MN 12010-I, MN 12011-I, MN 12012-I, MN 12013-I, MN 12014-I, MN 12015-I, MN 12016-I, MN 12017-I, MN 12018-I, MN 12019-I, MN 12020-I, MN 12021-I, MN 12022-I, MN 12023-I, MN 12024-I, MN12025-I, MN12026-I, and MN 12035-I. The specimens that were inside the palace and were later rescued are MN 10674-I, MN 10675-I, MN 10676-I, MN 10680-I, MN 10681-I, MN 10686-I, MN 10687-I, MN 10688, and MN 10690-I. Unfortunately, specimen MN 3323-I, which was previously studied by Derby (1895), was lost in the fire and is no longer available for examination.

The original accession numbers of specimens MN 10674-I, MN 10675-I, MN 10676-I, MN 10680-I, MN 10681-I, MN 10686-I, MN 10687-I, MN 10688-I, and MN 10690-I were inferred based on the positions of specimens in the collection before the fire. However, it is not possible to relate these individual specimens to their individual numbers. Because all of these specimens originated from the Monjolo dos Padres outcrop, we have designated the specimen in Figure 5.6 as MN 10674-I.

The samples deposited in the Universidade Federal de Mato Grosso collection include MP 533, MP 522a, b, MP 506, MP 507, MP 508 (part and counterpart), MP 510, and CD 310a-c.

We conducted a comparison of the material from the Paraná Basin with the hypotypes of Tropidoleptus carinatus (Conrad, 1839) studied by Williams (1913) from the Devonian of New York, United States of America. Currently, these samples are housed in the National Museum of Natural History/Smithsonian Institution collection with accession number USNM 59603 (1-88), originating from the Portage Formation. Additionally, Devonian specimens from the Amazonas Basin (Maecuru and Ererê formations) and Parnaíba Basin (Pimenteira and Cabecas formations) in Brazil were also analyzed. These samples are preserved in the collection of the Museu de Ciências da Terra/CPRM (DGM 2845-Ia-d; DGM 2902-I; DGM



**Figure 1.** Map of the Paraná Basin highlighting the Devonian outcrops. The red square indicates the region of the Juscimeira and Monjolo dos Padres outcrops, northwest–northwestern border of Paraná Basin, east–eastern border of Paraná Basin. Brazilian states: SC = Santa Catarina, PR = Paraná, SP = São Paulo, MG = Minas Gerais, MS = Mato Grosso do Sul, MT = Mato Grosso, GO = Goiás (adapted from Borghi and Fernandes, 2001; Scheffler et al., 2020).

2903-I), Instituto de Geociências/Universidade Federal do Rio de Janeiro (UFRJ 432-Bq, UFRJ 428-Bq), and Museu Nacional/UFRJ (MN 6379-I, MN 6382-I, MN 6391-I, MN 6398-I, MN 6400-I, MN 6402-I (not recovered after the fire).

For taxonomic identification, we consulted the Treatise on Invertebrate Paleontology (Harper, 2007) and various articles on

Tropidoleptus taxonomy across different Devonian regions worldwide.

Repositories and institutional abbreviations.—MN/UFRJ (MN-I), Museu Nacional of Universidade Federal do Rio de Janeiro, Rio de Janeiro, Brazil; IGEO/UFRJ (UFRJ-Bq),

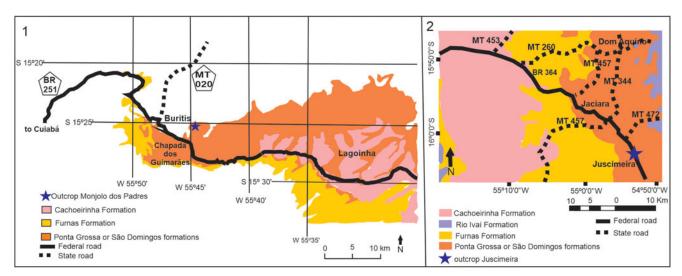
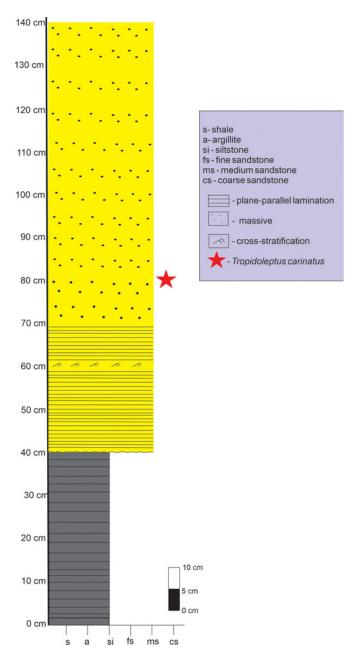


Figure 2. Location of the outcrops: (1) Simplified geological map of the Juscimeira outcrop region (Lacerda Filho, 2004). (2) Simplified geological map of the Monjolo dos Padres outcrop region (Grahn et al., 2010).



**Figure 3.** Sedimentological profile of the Juscimeira outcrop (adapted from personal correspondence, Henrique Z. Tomassi, 2021).

Instituto de Geociências of Universidade Federal do Rio de Janeiro, Rio de Janeiro, Brazil; UFMT (MP, CD), Universidade Federal de Mato Grosso, Cuiabá, Brazil; MCT/CPRM (DGM-I), Museu de Ciências da Terra of Serviço Geológico do Brasil, Rio de Janeiro, Brazil; NMNH/SI (USNM), National Museum of Natural History/Smithsonian Institution, Washington D.C., United States of America.

### Systematic paleontology

Class Rhynchonellata Williams et al., 1996 Order Tropidoleptida Schuchert, 1896 Superfamily Tropidoleptoidea Schuchert, 1896 Family Tropidoleptidae Schuchert, 1896 Genus *Tropidoleptus* Hall, 1857

Type species.—Strophomena carinata Conrad, 1839.

Tropidoleptus carinatus (Conrad, 1839) Figures 4, 5

- 1839 Strophomena carinata Conrad p. 64.
- 1857 Tropidoleptus carinatus (Conrad); Hall, p. 151.
- 1895 Tropidoleptus? sp. Derby, p. 76.
- 1897 Tropidoleptus rhenana Frech, p. 143, note 1, pl. 23a, fig. 9.
- 1902 *Tropidoleptus carinatus* var. *rhenana* (Frech); Drevermann, p. 99, pl. 12, figs. 11, 12.
- 1933 *Tropidoleptus carinatus* var. *maecurensis* Katzer; Clarke, p. 181, pl. 10, fig. 7a, b.
- 1947 Tropidoleptus sp. Caster, p. 126.
- 1967 *Tropidoleptus carinatus* (Conrad); Suárez-Riglos, p. 33, pl. 1, figs. 1–11.
- 1972 *Tropidoleptus carinatus* (Conrad); Carvalho, p. 25, pl. 2, figs. 1–14, pl. 3, fig. 1–3.
- 1977 *Tropidoleptus carinatus* (Conrad); Isaacson, pl. 1, figs. 20–37, pl. 2, figs. 1–3.
- 1977 *Tropidoleptus carinatus* (Conrad); Isaacson and Perry, p. 1115.
- 1977 *Tropidoleptus carinatus rhenana* (Frech); Isaacson and Perry, pl. 2, figs. 1–17.
- 1982 *Tropidoleptus platys* Cooper and Dutro, p. 43, pl. 4, figs. 13–21.
- 1987 *Tropidoleptus carinatus* (Conrad); Fonseca and Melo, pl. 1, figs. 1–9, pl. 2, figs. 1–12.
- 2004 Tropidoleptus sp. Racheboeuf et al., p. 104, fig. 4.10.
- 2010 Tropidoleptus carinatus (Conrad); Harper et al., figs. 1–6.

*Holotype*.—Undesignated holotype; type-series likely deposited in the American Museum of Natural History, United States of America (Gama Júnior, 2008; Queiroz et al., 2013).

Occurrence.—?Antarctica, unknown geological unit, Givetian (Harper et al., 2010); South Africa, Witterberg Group, Givetian-Frasnian (Boucot et al., 1983a; Penn-Clarke et al., 2018; Penn-Clarke, 2019); Germany, Lower Dark and Neichnerberg formations, Pragian-Emsian (Fuchs, 1982; Ubelacker et al., 2016); ?Australia, Cocktatoo Formation, Frasnian (Roberts, 1971); Belgium, unknown formation, Pragian–Emsian (Boucot et al., 1995); Bolivia, Huamampampa, Belén, and Sicasica formations, late Eifelian-?Frasnian (Isaacson, 1977, 1993; for age see Blieck et al., 1996; Troth et al., 2011); Brazil, São Domingos, Maecuru, Ererê, Pimenteira, and Cabeças formations, middle/late Eifelian-Givetian (Suárez-Riglos, 1967; Fonseca and Melo, 1987; Melo, 1988; this paper); Canada, Torbrook Formation, Lochkovian (Boucot, 1999); Colombia, Floresta and Gutierrez formations, ?Emsian-Frasnian (Caster, 1939; Morales, 1965; Barrett, 1986); Chile, Zorritas Formation, Eifelian or Givetian (Isaacson et al., 1985; Isaacson, 1993; Boucot et al., 1995, 2008); England, unknown formation, Pragian-Emsian (Boucot et al., 1995); Morocco, unknown

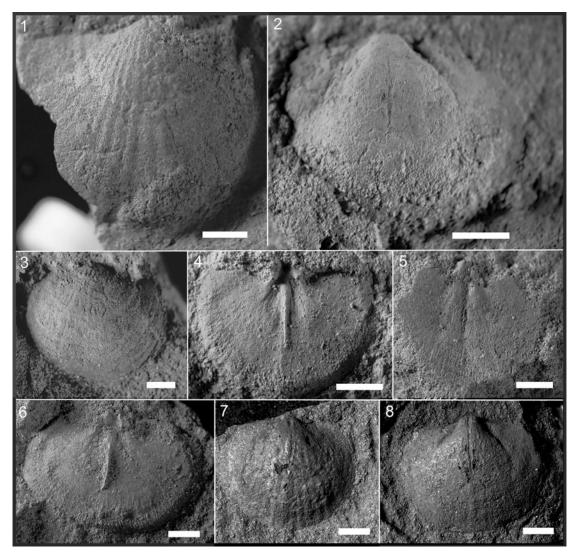


Figure 4. Tropidoleptus carinatus (Conrad, 1839): (1) exterior, ventral valve, MN 12003-I; (2) interior, ventral valve, MN 12005-I; (3) exterior, ventral valve, MN 12011-Ia; (4) interior, dorsal valve, MN 12012-I; (5) interior, dorsal valve, MN 12016-I; (6) interior, dorsal valve, MN 12014-I; (7) exterior ventral valve, MN 12015-Ib; (8) interior, ventral valve, MN 12015-Ia. All scale bars represent 5 mm.

formation, Pragian-Givetian (Jansen, 2001; Harper et al., 2010; Halamski and Balinski, 2013); France, unknown formation, Givetian (Rachebouef in Boucot et al., 1995); Central Sahara (Algeria and Niger), Tamesna Basin, Pragian-Givetian (Mergl and Massa, 2004); United States of America, Beechwood Limestone Member, Silica, "Beers Hill", Portage, Ludlowville and Mahantango formations, Givetian-Frasnian (Williams, 1913; Ellison, 1965; Kesling and Chilman, 1975; McGhee and Sutton, 1985); Spain, Salobral Formation, Givetian (Carls, 1988; Carls and Valenzuela-Ríos, 2002); Libya, Aouinet Ouenine II and Idri formations, ?Pragian-Frasnian (Freulon, 1964; Havlicek and Rohlich, 1987; Mergl and Massa, 1992); Mauritania, Aratane Formation, ?Pragian-Givetian (LeMaître, 1952; Sougy, 1964; Rachebouef et al., 2004); Czech Republic, Drakov Quartzite, Pragian (Isaacson and Chlupac, 1984); Peru, Cabanillas Formation, ?Eifelian (Newell, 1949; for age see Laubacher et al., 1982); Venezuela, Caño Grande and Caño del Oeste formations, ?Early Devonian-Givetian (Sanchez and Benedetto, 1983).

Description.—Length 10.5–25.2 mm, width 12.1–26.3 mm (Table 1, Fig. 6). The width is greater than the length, with the maximum width of valve found in median region; endopunctate; ~20 ribs arising at beak, ribs wider than the interspaces, relatively thick, and sparsely spaced, rarely multiplying by bifurcation; presence of concentric growth lines; contour of valves is subquadrate, semielliptical to semicircular, slightly concavo-convex to plano-convex. The beak is small, extending slightly beyond the hinge line, while the shell margins are rounded. Strophic, with a non-crenulate hinge line. The dorsal valve is almost planar, while the ventral valve is more concave; the median rib on ventral valve is slightly larger than others and extends to the anterior commissure.

Ventral interior.—Open delthyrium with a triangular shape, where height is smaller than the base size, and absence of deltidial plates; prominent teeth are present and diverge anterolaterally; muscular field broad, long, and weakly imprinted, with an inflated umbonal region; presence of delicate myophragm.

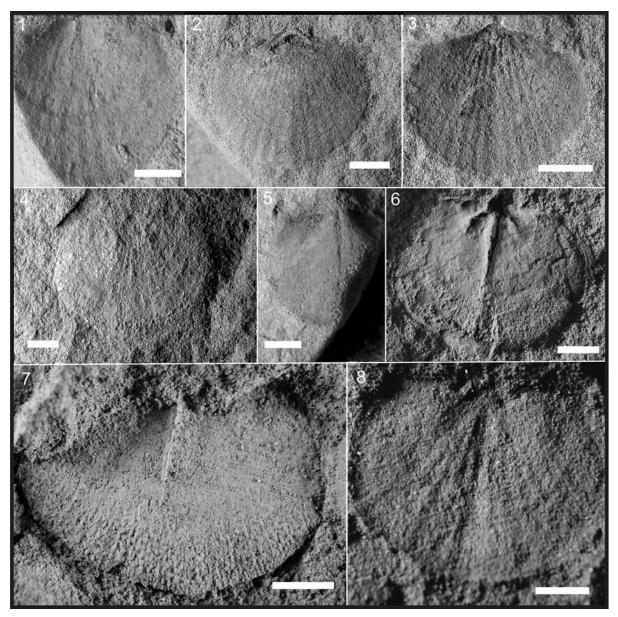


Figure 5. Tropidoleptus carinatus (Conrad, 1839): (1) exterior, ventral valve, MP 533; (2) exterior, dorsal valve, MP 522a; (3) exterior, ventral valve, MP 522b; (4) exterior, dorsal valve, CD 310b; (5) interior, ventral valve, MP 507; (6) interior, dorsal valve, MN 12017-I; (7) interior, MN 12004-Ia; (8) interior, dorsal valve, MN 12022-I. All scale bars represent 5 mm.

Dorsal interior.—Well-excavated dental pits are delimited by lateral lobes of the notothyrial platform, poorly preserved cardinal process projects above the hinge line. Crura extend anterolaterally as descending lamellae from the lateral lobes of the notothyrial platform. A thin myophragm starts from the anteromedial region of the notothyrial platform and extends over more than half of the valve, thinning anteriorly.

Material.—MN 10674-I, MN 10675-I, MN 10676-I, MN 10680-I, MN 10681-I, MN 10686-I, MN 10687-I, MN 10688-I, MN 12002-Ia,b, MN 12003-Ia,b, MN 12004-Ia,b, MN 12005-I, MN 12006-I, MN 12007-I, MN 12008-I, MN 12009-I, MN 12010-I, MN 12011-I, MN 12011-I, MN 12013-I, MN 12014-I, MN 12015-Ia,b, MN 12016-I, MN 12017-I, MN 12018-I, MN 12019-I, MN 12020-I, MN 12021-Ia,b, MN 12022-I, MN

12023-Ia, b, MN 12024-I, MN 12025-I, MN 12026-Ia-e, MN 12035-Ia-c, MP 533, MP 522a, b, MP 506, MP 507, MP 508 (part and counterpart), MP 510, CD 310a-c (55 samples).

*Provenance.*—Outcrops Monjolo dos Padres and Juscimeira, São Domingos Formation (possibly Givetian).

Observation.—Tropidoleptus reported by Kayser (1897) from the Los Espejos Formation in Argentina was synonymized as Castellaroina (see Boucot in Amos, 1972; Harper and Boucot, 1978; and Benedetto and Montoya, 2015). Tropidoleptus cited by Thomas (1905) in Argentina may not belong to the genus either.

Remarks.—The specimens analyzed in this study exhibit the main morphological features of *Tropidoleptus*, as described by

Table 1. Measurements of the analyzed specimens.

Specimen	Length (mm)	Width (mm)	L/W
MN 12003-Ia	18.2	22.3	0.82
MN 12024-I	13.3	_	_
MN 12004-Ia	13.1	18.9	0.69
MN 12004-Ib	12.8	22.3	0.57
MN 12005-I	18.0	24.1	0.75
MN 12006-I	9.2	13.5	0.68
MN 12009-I	25.2	30.0	0.84
MN 12035-Ia	19.0	25.9	0.73
MN 12035-Ib	14.9	19.2	0.78
MN 12010-I	20.5	26.3	0.78
MN 12011-I	15.6	19.6	0.79
MN 12012-I	16.8	24.0	0.7
MN 12013-I	10.5	13.3	0.79
MN 12016-I	16.1	18.4	0.87
MN 12014-I	20.2	29.3	0.69
MN 12019-I	14.4	16.4	0.87
MN 12020-I	21	27.8	0.75
MN 12015-Ia	18.7	22.3	0.81
MN 12015-Ib	14.5	18.6	0.78
MN 12021-Ia	15.7	20.9	0.75
MN 12021-Ib	14.9	19.7	0.76
MN 12022-I	22.6	32.2	0.70
MN 12017-I	21.1	27.6	0.76
MN 12026-Ia	23.9	33.4	0.72
MN 10674-I	13.6	_	_
MP 522a	11.7	15.2	0.77
MP 522b	10.8	13.0	0.83
MP 507	12.5	_	_
MP 508	10.5	_	_
MP 510	10.7	12.1	0.88
MP 533	12.3	_	_
CD 310a	12.8	_	_
CD 310b	18.5	23.0	0.8
Mean	15.9	21.7	0.76
Standard Deviation	4.14	5.75	0.07

Conrad (1839), Isaacson and Perry (1977), Harper (2007), and Harper et al. (2010). These features include: (1) endopunctate shell surface; (2) subquadrate, semielliptical to semicircular

contour of valves, slightly concavo-convex to plano-convex; (3) slightly larger median rib on ventral valve compared to other ribs; (4) ribs originating at the beak, wider than interspaces, relatively thick, and sparsely spaced, rarely multiplying by bifurcation; (5) presence of concentric growth lines; and (6) crura extending anterolaterally as descending lamellae.

Based on these observations, there is no doubt that *Tropidoleptus* was present in the Devonian of the Paraná Basin. We concur with the inclusion of *Tropidoleptus* in the order Tropidoleptida, as discussed by Harper et al. (2010).

Several species have been proposed within the genus *Tro-pidoleptus*. Other than the type species, at least four additional species can be recognized: *T. occidens* (Hall, 1860); *T. platys* Cooper and Dutro, 1982; *T. fascifer* Kayser, 1897; and *T. rhenanus* (Frech, 1897).

Tropidoleptus fascifer was proposed by Kayser (1897) from Argentina. However, it was subsequently designated as the type species of the subgenus *Castellaroina* by Boucot in Amos (1972), which was later elevated to genus rank by Harper and Boucot (1978). Benedetto and Montoya (2015) proposed an emended diagnosis for the genus *Castellaroina* and provided a detailed description of the species *Castellaroina fascifer* (Kayser, 1897). As a result, *T. fascifer* is no longer considered a valid species of *Tropidoleptus*.

Tropidoleptus rhenanus (= rhenana) was proposed by Frech (1897) based on species from the Devonian of Germany. However, Drevermann (1902) conducted a comprehensive analysis of extensive German, American, and Bolivian material and concluded that *T. rhenana* could not be considered a separate species from *T. carinatus*. The external morphological characteristics (e.g., the number of ribs, length/width ratio, and median rib) are similar among these specimens. Drevermann (1902)

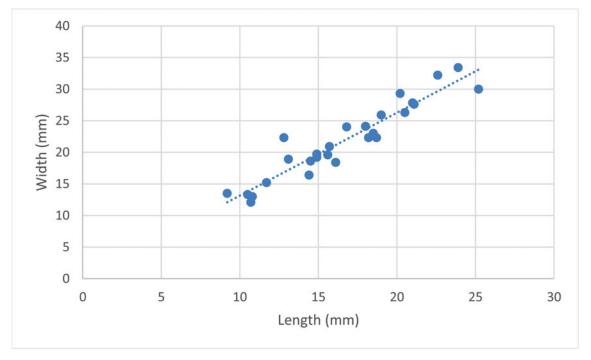


Figure 6. Plots of measurements of length versus width of analyzed *Tropidoleptus carinatus* (Conrad, 1839); n = 27.

suggested a possible distinction based on the extent of the muscular field, which extends beyond half of the valve. However, it has been shown in figures by Isaacson and Perry (1977) that *T. carinatus* also can exhibit these features. Upon careful comparison, it is apparent that "*T. rhenanus*" has slightly more angular ribs compared to *T. carinatus* (see Isaacson and Perry, 1977; Harper et al., 2010; Schemm-Gregory and Henriques, 2013). This suggests that "*T. rhenanus*" can be classified as a subspecies (*Tropidoleptus carinatus rhenanus*), rather than a distinct species, which aligns with the findings of Drevermann (1902) and Isaacson and Perry (1977).

Tropidoleptus occidens was proposed by Hall (1860) for specimens from the Devonian of Iowa, United States of America. However, the limited number of described specimens and their poor preservation make it challenging to conduct a thorough comparison with *T. carinatus*. One notable difference is the presence of two enlarged median ribs on the ventral valve, as described by Hall (1860), which is not a typical characteristic of *Tropidoleptus*. It should be noted that the illustrations by Hall and Clarke (1894) do not provide a clear characterization of these two enlarged median ribs. In contrast, *Tropidoleptus* typically possesses only one enlarged median rib (see Harper, 2007; Harper et al., 2010). Stainbrook (1938) synonymized *T. occidens* as *Stropheodonta occidens* (Hall, 1860).

Tropidoleptus platys was proposed by Cooper and Dutro (1982) based on specimens from the Devonian of New Mexico, United States of America. However, the characteristics identified by these authors are not sufficiently robust to designate T. platys as a distinct species for the following reasons. (1) The obsolescence of ribs on the anterior margin noted in T. platys may be attributed to taphonomic factors rather than being a consistent characteristic across all specimens (some specimens of *T. carinatus* also exhibit this feature, as indicated by Cooper and Dutro, 1982; see also Isaacson and Perry, 1977, for an example. (2) Cooper and Dutro (1982) mentioned that specimens of T. carinatus from the Kashong Shale are larger than T. platys, while T. carinatus specimens from the Tully Limestone are smaller than T. platys. Hence, size cannot be considered a diagnostic feature in this case. The samples of T. carinatus described by Isaacson and Perry (1977) exhibit lengths of 15-30 mm, while Isaacson (1977) recognized a length range of 3.4–20.4 mm in Bolivian specimens. The specimens described by Cooper and Dutro (1982) have a lengths of 13.9–16.7 mm, which is within the size range expected for T. carinatus (3) The supposed smaller concavity and less-pronounced costate in T. platys compared to T. carinatus are not evident when comparing the illustrations by Cooper and Dutro (1982) with the specimens described by us and in other papers (e.g., Williams, 1913; Isaacson, 1977; Isaacson and Perry, 1977; Fonseca and Melo, 1987; Harper et al., 2010). (4) The more transverse shell of T. rhenanus (here considered synonymous with T. carinatus) compared to T. platys is not a robust diagnostic character because it is highly susceptible to taphonomic influences. Therefore, based on the characteristics presented here, T. platys is considered synonymous with T. carinatus.

A comparison of *T. carinatus* specimens from the Paraná, Parnaíba, and Amazonas basins, as well as the hypotypes from New York (hypotypes of Williams, 1913), revealed no significant differences among them. The observed variations were consistent with intraspecific variations, as discussed by Isaacson and Perry (1977), confirming the presence of the same species across these regions (Fig. 7). Nevertheless, two subtle main differences were identified in the analyzed Brazilian specimens compared to those described by Isaacson and Perry (1977): (1) the ventral myophragm appears to be less vigorous, and (2) the median septum (referred to as the myophragm here) is less developed. These differences also were recognized by Fonseca and Melo (1987) in Devonian specimens from the Parnaíba Basin.

Finally, in the existing literature, several subspecies of *Tropidoleptus carinatus* have been proposed, including *T. carinatus freuloni* Boucot, Massa, and Perry, 1983b; *T. carinatus nigeriensis* Boucot, Massa, and Perry, 1983b; and *T. carinatus titanius* Mergl and Massa, 2004. The specimens studied here can be clearly distinguished from *T. carinatus titanius* by their smaller size (with a maximum width of ~33 mm, compared to 56 mm in *T. carinatus titanius*). However, in order to reliably differentiate the *T. carinatus specimens* studied here from *T. carinatus freuloni* and *T. carinatus nigeriensis*, it would be necessary to examine specimens the latter two subspecies. Analysis of the photos presented by Boucot et al. (1983b) was inconclusive for this purpose.

# Paleoenvironmental and paleobiogeographic considerations

Paleoenvironment.—Tropidoleptus carinatus (Conrad, 1839) is a species commonly associated with shoreface environments worldwide, although it also could occur in shallower areas of the transitional offshore (see Isaacson and Perry, 1977; Boucot et al., 1995; Harper et al., 2010). The Paraná Basin specimens of Tropidoleptus carinatus are found in friable, well-sorted sandstones. In the Juscimeira outcrop, the level containing Tropidoleptus is characterized by massive medium-grained sandstone (personal correspondence, Henrique Z. Tomassi, 2021). Specimens in the Monjolo dos Padres outcrop were found in sandstones with hummocky cross-stratification (tempestites), interbedded with shales (Quadros and Melo, 1989). Most specimens are disarticulated, and some show signs of abrasion, all of which may be a result of the naturally abrasive nature of the lithology (friable sandstone) and the high-energy paleoenvironment, rather than indicating long-distance transport. Tropidoleptus carinatus was not found in thinner facies such as argillite, siltstone, or shale, suggesting that this species was also restricted to shallower environments in the Paraná Basin, likely between the distal shoreface and proximal offshore. Further taphonomic studies are required to provide additional support for this interpretation.

Paleobiogeography.—Tropidoleptus is another extra-Malvinokaffric brachiopod genus found in the Devonian deposits of the Paraná Basin, alongside other genera such as Chonostrophia?, Cryptonella, Plicoplasia?, Pustulatia?, and Schellwienella (Clarke, 1913; Cerri, 2013; Penn-Clarke and Harper, 2021; Rezende and Isaacson, 2021; Sedorko et al., 2021; Videira-Santos et al., 2022).

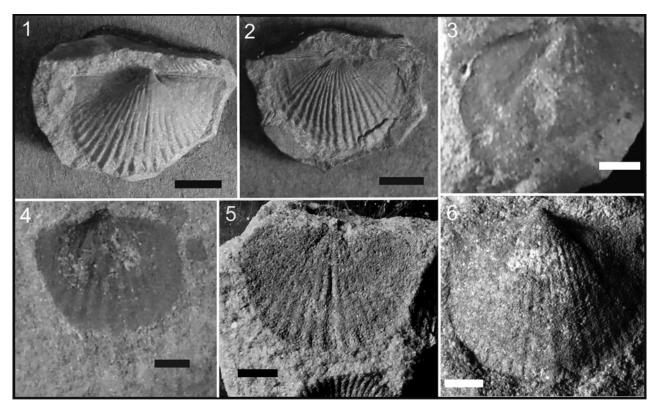


Figure 7. Tropidoleptus carinatus (Conrad, 1839). (1) Exterior, ventral valve, Harford Mills, New York, USA, Portage Formation, USNM 59603 (21); (2) exterior, dorsal valve, Harford Mills, New York, USA, Portage Formation, USNM 59603 (88); (3) interior, ventral valve, BR-020, Itainópolis, Piauí, Brazil, Pimenteira Formation, MN 6391-I; (4) exterior, ventral valve, BR-020, Itainópolis, Piauí, Brazil, Pimenteira Formation, MN 6400-I; (5) interior, dorsal valve, Serra de Ererê, Monte Alegre, Pará, Brazil, Ererê Formation, DGM 2845-Ia; (6) exterior, ventral valve, Maecuru River, Monte Alegre, Pará, Brazil, Maecuru Formation, DGM 2902-I. All scale bars represent 5 mm.

The genus *Tropidoleptus* originated in the Lochkovian of Nova Scotia (Canada), part of the ancient Avalonia/Baltica region (see Boucot, 1999; Robardet, 2003), belonging to the "Old World Realm" (Dowding and Ebach, 2018; Dowding et al., 2022). From its origin, *Tropidoleptus* dispersed to different parts of the world, including Europe, North Africa, and South America, the central portion of North America, South Africa, and possibly Antarctica, indicating its cosmopolitan distribution (Isaacson and Perry, 1977; Harper, 2007; Harper et al., 2010) (Fig. 8). However, the genus eventually became extinct during the Late Devonian extinction event, as noted by Harper et al. (2010).

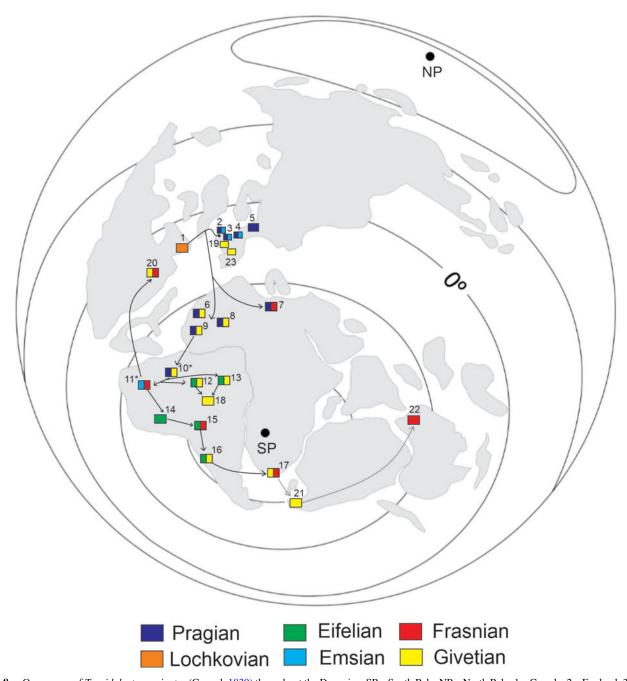
According to paleogeographical reconstructions, Avalonia/ Baltica and the rest of Europe were closely positioned during the Early Devonian, with only a narrow Rheic Ocean separating them (Robardet, 2003). Additionally, southeastern Europe, including the Armorican Massif, the Iberian Peninsula, and possibly central Europe (Bohemia) and perhaps the Moesian Terrane, were positioned adjacent to the Gondwana plate, forming its northern border (Plusquellec et al., 1997; Robardet, 2003; Vaida et al., 2005). This paleogeographic context suggests a potential pathway for the rapid migration of *Tropidoleptus* from Avalonia/Baltica to Bohemia and northern Gondwana during the Early Devonian (Fig. 8).

During the Eifelian global transgression (see Haq and Schutter, 2008), the sea advanced into the Amazonas and Parnaíba basins, which are part of the Amazonian sub-region, as defined by Penn-Clarke and Harper (2021). *Tropidoleptus* 

along with other brachiopods, echinoderms, trilobites, and various taxa, migrated from regions such as Armorica, Bohemia, and North Africa towards the northern (Amazonas Basin) and northeastern (Parnaíba Basin) parts of Brazil between the Emsian and Eifelian (Fonseca and Melo, 1987; Díaz-Martínez et al., 1993; Scheffler et al., 2011; Carvalho and Ponciano, 2015; Fonseca, 2015; Scheffler, 2015). This migration route highlights the dispersal of organisms during this period of time.

Penn-Clarke and Harper (2021) reached the conclusion, based solely on brachiopods, that the Amazon Basin predominantly hosted taxa from the eastern Americas, including Malvino-kaffric elements, while the Parnaíba Basin exhibited a combination of eastern Americas and Malvinokaffric taxa. Nevertheless, as mentioned earlier, it is important to recognize that the faunal composition of these basins also was influenced by elements of the Old World.

During the Eifelian, *Tropidoleptus* also extended its distribution to various other regions of present-day South America (e.g., Bolivia, ?Chile, and ?Peru). The influence of the Northern Gondwana fauna, as noted earlier during the Middle Devonian in South America, has been documented by several studies (e.g., Isaacson and Sablock, 1988; Díaz Martínez et al., 1993; Racheboeuf et al., 1993; Blieck et al., 1996; Tourneur et al., 2000). The evidence presented in this study suggests that the faunal distribution in southwest Gondwana was influenced by sea-level events and facilitated dispersal, which in turn



**Figure 8.** Occurrence of *Tropidoleptus carinatus* (Conrad, 1839) throughout the Devonian. SP = South Pole, NP = North Pole; 1 = Canada, 2 = England, 3 = Belgium, 4 = Germany, 5 = Czech Republic, 6 = Morocco, 7 = Libya, 8 = Central Sahara (Algeria/Niger), 9 = Mauritania, 10 = Venezuela, 11 = Colombia, 12 = Amazonas Basin (Brazil), 13 = Parnaiba Basin (Brazil), 14 = Peru, 15 = Bolivia, 16 = Chile, 17 = South Africa, 18 = Parnai Basin (Brazil), 19 = France, 20 = United States of America, 21 = Antarctica, 22 = Australia, 23 = Spain. Dotted arrow = migration with doubts (adapted from Habicht, 1979; Raymond, 1987; Melo, 1988; Golonka, 1994; Scotese et al., 1999; di Pasquo et al., 2009; Boucot et al., 2013; Ribeiro, 2020). \*There are doubts about the Pragian age for Venezuela and the Emsian age for Colombia; *Tropidoleptus carinatus* possibly entered these locations later.

influenced the distribution and mixing patterns between various paleobiogeographic sub-regions, as previously postulated by several authors (e.g., Melo, 1988; Dowding et al., 2022). Additionally, differences in latitudes have an effect on taxonomic distribution and bioregionalization, as supported by the findings of Penn-Clarke and Harper (2021).

According to Penn-Clarke and Harper (2021) and Dowding et al. (2022), further biogeographic studies are necessary to

reassess the status of sub-regions within the Malvinokaffric Realm. However, these studies should not rely solely on one phylum of organisms. In the case of southwest Gondwana, the similarity between taxa from different Malvinokaffric basins appears to vary depending on the phylum being considered. For instance, the work of Scheffler et al. (2013) highlights the lack of similarity between crinoids from the Paraná Basin and those from other basins in South America and South Africa,

illustrating the need for a comprehensive examination across multiple phyla.

Tropidoleptus, known for inhabiting warmer waters, traditionally has been regarded as an end marker of the typical endemic fauna of certain Devonian paleobiogeographic realms (Dowding et al., 2022). However, its presence has been documented in these areas even prior to the collapse of the aforementioned paleobiogeographic realms (Isaacson, 1977; Penn-Clarke et al., 2018; Penn-Clarke, 2019). For instance, the Malvinokaffric Realm is thought to have ceased to exist by the Givetian (Boucot, 1988; Li, 2010; Dowding and Ebach, 2018; Penn-Clarke and Harper, 2021; Rezende et al., 2021). Nonetheless, there are records of *Tropidoleptus* in Malvinokaffric regions (Bolivia and South Africa) in older deposits (e.g., Isaacson, 1977; Harper et al., 2010; Troth et al., 2011; Penn-Clarke et al., 2018; Penn-Clarke, 2019).

The arrival of *Tropidoleptus* in the Paraná Basin during the Givetian coincides with the decline of the typical Malvinokaffric fauna that had previously inhabited the basin, as well as with the occurrence of a "Lilliput effect" in the relict fauna (Bosetti et al., 2010; Grahn et al., 2016). The occurrence of *T. carinatus* in strata of presumed Givetian age in the Paraná Basin supports the idea that the Parnaíba, Amazonas, and Paraná basins had been interconnected since the early Givetian, and that warm waters entered the Paraná Basin, which was located near the south pole at that time (Grahn et al., 2016). This coincided with the Acadian orogeny and the circulation of warm water counterclockwise currents associated with a transgression and the northward movement of western Gondwana, facilitating the influx of these warm currents (Isaacson and Sablock, 1990; Grahn et al., 2016).

It is likely that *Tropidoleptus* migrated to the Paraná Basin from the Parnaíba and/or Amazonas basins during this transgression episode, supporting the idea that elements of the boreal fauna could colonize southern environments during the Devonian. This migration of invertebrates from the Parnaíba and Amazonas basins to the Paraná Basin in the Eifelian/Givetian transition, associated with the transgression of warmer waters, also is observed in echinoderms (e.g., Scheffler et al., 2011, 2017; Scheffler, 2015).

### **Conclusions**

In this study, we have provided confirmation of the presence of *Tropidoleptus carinatus* (Conrad, 1839) in Devonian strata of the Paraná Basin. The morphological characteristics of the specimens analyzed align with the accepted variations for this species that have been described in the literature. The migration of *T. carinatus* is thought to have occurred during a late Eifelian/early Givetian marine transgression, which facilitated the influx of warmer waters from the Amazonas and Parnaíba basins into the Paraná Basin.

The presence of *T. carinatus* can serve as an indicator of shoreface and shallower portions of the transitional offshore environments, when found in situ. In the context of the Paraná Basin, its occurrence can be used to suggest that the associated strata are not older than the late Eifelian. However, in order to provide further support to these interpretations, it is essential to

conduct a taphonomic study and utilize dating methods other than fossil composition to accurately determine the age of the outcrop where *T. carinatus* is found. Such comprehensive investigations can provide valuable insights into the paleoenvironmental context and temporal distribution of this species within the basin.

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### **Declaration of competing interests**

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

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