EFFECT OF SiO₂/Al₂O₃ RATIO ON THE THERMAL **REACTIONS OF ALLOPHANE**

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Abstract--Differences were found in the differential thermal analysis curves and in the temperatures of new-phase development between allophanes of high $(1.91-1.99)$ and low $(1.47-1.53)$ SiO₂/Al₂O₃ ratios. The endothermic peak due to continuous dehydration and dehydroxylation was at higher temperatures (153[°]-185°C) for allophanes with high SiO_2/Al_2O_3 ratios and at lower temperatures (148°–165°C) for those with low $SiO₂/Al₂O₃$ ratios. The temperature of the exothermic peak was lower and the height affected more by the exchangeable cation content for allophanes with high ratios than for those with low ratios. New phases did not develop in allophanes having high SiO_2/Al_2O_3 ratios even after they were heated to 1000°C, above the temperature of the exotbermic peak. In contrast, a symptomatic development of new phases was noted in allophanes with low SiO_2/Al_2O_3 ratios at 900°C, below the temperature of the exothermic peak. The effect of SiO_2/Al_2O_3 ratio in the thermal behavior of allophane strongly suggests that differences in the structure are closely associated with the chemical composition of this material.

Key Words--Allophane, Dehydration, Dehydroxylation, DTA, Silica/alumina ratio, Thermal reactions.

INTRODUCTION

In differential thermal analysis (DTA), allophane generally shows a broad endothermic peak between 100° and 300° C and an exothermic peak between 800° and 1000° C due to continuous dehydration and dehydroxylation, and to the formation or nucleation of mullite and/or gamma-alumina, respectively (Mitchell *et al.,* 1964; Wada and Harward, 1974; Fieldes and Claridge, 1975; Wada, 1977). The DTA curve has been shown to be susceptible to various factors. Thus, Fieldes (1955) differentiated "Allophane B" from "Allophane A" in New Zealand volcanic ash soils on the basis of the lack, or appearance of an exothermic peak near 900°C. Fieldes (1957) further mentioned that the endothermic peak of "Allophane B" tended to be lower in temperature than that of "Allophane A." Miyauchi and Aomine (1966) indicated that the exothermic peak of allophane varied in temperature, height, and area with exchangeable cation. The exothermic reaction also was observed to be affected by the method of separation or pretreatment of the material (Miyazawa, 1966; Campbell *et al.,* 1968). For amorphous silicoaluminas synthesized as analogues of allophane, the shape of the exothermic peak was correlated with their $SiO₂/Al₂O₃$ ratios (Wada, 1977). As allophanes have a wide range of chemical composition, $(SiO₂/Al₂O₃ = 1.0$ to 2.0; Yoshinaga, 1966; Wada and Yoshinaga, 1969; Henmi and Wada, 1976), their thermal behavior can be expected to be influenced by the chemical composition.

The present investigation is concerned with the effect of the $SiO_2/A1_2O_3$ ratio on the DTA curve of allophane and on the temperature at which new phases develop.

MATERIALS AND METHODS

Fine clay fractions ($< 0.2 \mu$ m) were separated by centrifugation from weathered volcanic ash and pumice

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from Japan (Table 1). Before the separation, the ash and pumice were repeatedly treated with H_2O_2 to remove organic matter, gently pulverized, and dispersed by sonification at 20 kHz. X-ray powder diffraction analysis (XRD) and infrared spectroscopy (IR) showed that the fine clay separates were typical allophanes. Electron optical observation affirmed that the clays were virtually "pure" allophane and that they appeared as various size aggregates with irregular shape. Many small particles, probably consisting of hollow spherules in three dimensions, were seen as ring shapes with diameters of 35 to 50 A in electron micrographs (see also, Kitagawa, 1971; Wada *et al.,* 1972; Henmi and Wada, 1976; Wada and Wada, 1977; Shoji and Saigusa, 1977). Only small amounts of imogolite fibers (Yoshinaga *et al.,* 1968; Wada *et al.,* 1970) were recognized in the clays; imogolite is commonly admixed in allophanes from weathered volcanic ash and pumice (Henmi and Wada, 1976). The clays were treated with sodium dithionite-citrate-bicarbonate (Mehra and Jackson, 1960) and 2% Na₂CO₃ (Jackson, 1956a) and washed with 1 N NaCl solution to exchange the clay with Na. The Na-saturated product was dialyzed against distilled water until the solution was free of CI ion. H-clay was prepared from the Na-clay by treating it with a H-form ion-exchange resin (IR-120). Part of the Na-clay was dissolved completely with 0.15 M sodium oxalate-oxalic acid mixture (pH 3.5) at 90 $^{\circ}$ C (Higashi and Ikeda, 1974; Henmi and Wada, 1976; Wada and Wada, 1976). The Si and A1 in the resulting solution were determined using an atomic absorption spectrophotometer (Hitachi 508) by the method of Henmi and Wada (1976).

DTA was carried out using an automatic recording instrument with chromel-alumel thermocouples (Rigaku Thermoflex 8001 . The sample (100 mg) was losely packed in a platinum holder after being freeze-dried and

1 Henmi (1977).

² Molar ratio of the $<$ 0.2- μ m fraction treated with sodium dithionite-citrate-bicarbonate and 2% Na₂CO₃.

stored over a saturated solution of $Mg(NO₃)₂·6H₂O$, according to the method of Mackenzie and Mitchell (1970). Calcined Al_2O_3 was used as a reference material. A heating rate of 20° C/min was employed in a static atmosphere (air). Silver nitrate standard was used to calibrate the furnace temperature of the instrument (Jackson, 1956b). The temperatures of the endothermic and exothermic peaks were estimated from the positions of the top or bottom of the peak on the chart.

XRD was carried out using $CuKa$ radiation (30 kV, 10 mA) for the freeze-dried clays after they had been heated for 5 hr to 1200° C at 100° C increments to determine the temperature at which mullite formed.

RESULTS AND DISCUSSION

The samples listed in Table 1 range in SiO_2/Al_2O_3 ratios from 1.47 to 1.99 and fall into two groups with low (-1.5) and high (-1.9) SiO₂/Al₂O₃ ratios. The DTA curves are characteristic of allophane and show a broad endotherm and an exotherm in the temperature ranges $150^{\circ}-187^{\circ}$ C and $946^{\circ}-977^{\circ}$ C, respectively. Figure 1 shows the temperature of endothermic and exothermic peaks plotted against the $SiO₂/Al₂O₃$ ratios of the clays. Evidently these temperatures are influenced by the SiO_2/Al_2O_3 ratios as well as by the nature of the exchangeable cation.

Endotherrnic peak

The endothermic peak temperatures are always higher for H-clays than for Na-clays. Higher endothermic peak temperatures appear in both the H- and Na-clays with high $SiO₂/Al₂O₃$ ratios, but the increase is smaller for Na-clays. The order with decreasing temperature is as follows: H-clay with high SiO_2/Al_2O_3 ra $tios > H-clay$ with low ratios $> Na-clay$ with high ratios > Na-clay with low ratios. This order corresponds to the surface acid strength of allophane (Henmi, 1977) because with stronger surface acidity, water is probably more strongly adsorbed as a base on the allophane surface. It is inferred that water and hydroxyls are more strongly bonded in allophane with high SiO_2/Al_2O_3 ratios than in those with low ratios. Barshad (1965) showed that the temperature of the endotherm due to dehydration of montmorillonite and vermiculite varied

Figure 1. Effect of $SiO₂/Al₂O₃$ ratio on the temperature of endothermic and exothermic peaks for H- and Na-exchanged allophanes.

with exchangeable cation. The temperature variation for montmorillonite has been attributed to differences in the hydration energies of different cations (Mackenzie, 1970). Therefore, the difference in endothermic peak temperature between the H- and Na-allophanes can be explained by the different hydration energies of the cations.

Exothermic peak

The temperature of the exothermic peak for allophanes with high SiO_2/Al_2O_3 ratios is higher for H-clays than for Na-clays. This is consistent with the results of Miyauchi and Aomine (1966). As seen in Figure 1, however, exothermic peak temperatures are almost the same for H- and Na-clays having low $SiO₂/Al₂O₃$ ratios. The temperature of the exothermic reaction is also dependent on the chemical composition, if allophane is treated with NaCI instead of a H-ion exchange resin. H-ion saturation results in a nearly constant temperature (\sim 975°C) for the allophane exotherm, irrespective of its $SiO₂/Al₂O₃$ ratio.

| Clay | Na-clay I _{Na} | H-clay I, | I_H/I_{Na} | |
|------|----------------------------|--------------|--------------|--|
| Ki-P | 42 | 43 | 1.02 | |
| Ky-P | 51 | 90 | 1.76 | |
| PA-P | 15 | 118 | 7.83 | |
| VA | 35 | 108 | 3.09 | |

Table 2. Exothermic peak height of H- and Na-allophanes.'

¹ In terms of Δ e.m.f. (μ V). I = peak intensity.

The exothermic peak heights listed in Table 2 in terms of differential electromotive force $(\Delta e.m.f.)$ vary with SiO_2/Al_2O_3 ratio and with exchangeable cation. Stronger peaks (42–51 μ V) are associated with Naclays having low $SiO₂/Al₂O₃$ ratios than those having high ratios (15-35 μ V). The intensity of the exothermic peak was correlated with $SiO₂/Al₂O₃$ ratio for allophane from a stream-bed deposit below Silica Spring outlet on Mt. Ruapehu of New Zealand by Wells *et al.* (1977) who also reported that the peak decreased in sharpness (intensity) with an increase in the $SiO₂/Al₂O₃$ ratio. This observation was also made for the Na-clays studied in the present investigation. On the other hand, stronger exothermic peaks (108-118 μ V) were noted for H-clays with high $SiO₃/Al₃O₃$ ratios than for those with low ratios (43–90 μ V). Thus, exchanging Na⁺ for H⁺ causes a considerable reduction in the intensity of the exothermic peak for all samples (Table 2). Miyauchi and Aomine (1966) also showed that the height of exothermic peak was greater for allophane saturated with H^+ than with $Na⁺$. The ratio of the exothermic peak height for H-clays to that for Na-clays differs from one clay to another, indicating that among allophanes of different $SiO₂/Al₂O₃$ ratios, peak height enhancement is more pronounced for clays with high $SiO₂/Al₂O₃$ ratios (3.09 to 7.83) than for those with low ratios (1.02 to 1.76).

New-phase-development

The correlation between the chemical composition and the temperature at which new phases develop is given in Table 3. Only data for H-clays are compared because the temperature of exothermic peak of H-clays is essentially unaffected by the $SiO₂/Al₂O₃$ ratio, although the peak heights vary with the ratios (Figure 1, Table 2). The temperature at which mullite was detected by XRD is regarded as the new-phase-development temperature. This temperature is different for clays with high $SiO₉/Al₂O₃$ ratios and with low ratios. New phases developed in all the clays after heating to 1200° C. Although the temperature of the exothermic peak of H-clays is between 970° and 977° C, new phases were not detected in clays with high SiO_2/Al_2O_3 ratios until they had been heated to $>1000^{\circ}$ C. However, a symptomatic development of new phases was noted in clays with low ratios at temperatures below the exothermic peak. The results indicate that crystallization

Table 3. Detection of mullite in heated fine clays by X-ray powder diffraction analysis.

| Clav ¹ | | | Temperature (°C) | | |
|-------------------|-----|-----|------------------|------|------|
| | 400 | 800 | 900 | 1000 | 1200 |
| Ki-P | | | $\,{}^+$ | | |
| Ky-P | | | + | | |
| $PA-P$ | | | | | |
| VA | | | | | |

¹ H-clay. $+$ = detected, \pm = symptomatic, $-$ = not detected.

to mullite is easier from allophanes with low $SiO₂/Al₂O₃$ ratios than from those with high ratios, wherein a $SiO₂$ component may inhibit the crystallization of mullite or growth of its crystal size. Ossaka (1962) noted that mullite crystallized as low as 840°C, before the exothermic peak appeared, from two materials having $SiO₂/Al₂O₃$ ratios of 0.14 and 1.04, which were referred to as "precipitated allophane" in view of their occurrence as stream deposits (Ossaka, 1960, 1961).

Interpretation

The experimental results discussed above indicate that the $SiO₂/Al₂O₃$ ratio influences the thermal behavior of allophane and suggest that structural differences are closely associated with chemical composition. Henmi and Wada (1976) noted differences in the amount of Al in 4-fold coordination (A^{IV}) among allophanes of different chemical composition, and other workers have assigned A1 to both 4- and 6-fold coordinations on the basis of X-ray fluorescence spectroscopic results (Egawa, 1964; Udagawa *et aL,* 1969; Okada *et al.,* 1975). Henmi and Wada (1976) noted that with higher SiO_2/Al_2O_3 ratio, more Al^{IV} is present, leading to the expectation that the negative charge due to the presence of tetrahedral AI increases with an increase in the ratio. In fact, the cation-exchange capacity (CEC) measured at pH 7 is higher for allophanes with high $SiO₂/Al₂O₃$ ratios (~1.9) than those with low ratios (-1.3) (unpublished data from this laboratory). It is, therefore, reasonable to suggest that the variation in CEC with chemical composition is partly responsible for the differences in thermal behavior among allophanes of different SiO_2/Al_2O_3 ratios. The effects of the nature of the exchangeable cation on the DTA curve are greater for allophanes with high SiO_2/Al_2O_3 ratios than those with low ratios (Figure l, Table 2). This fact can be explained by the higher CEC, and therefore, the greater amount of exchangeable cations in allophanes with high $SiO₂/Al₂O₃$ ratios compared with those with low ratios.

Brindley and Fancher (1969) proposed that allophane has a defect kaolinite structure wherein O is replaced by OH, thereby placing H on the outer edge of the layer, or by replacing Si by 4H so as to produce vacancies.

The variation in thermal behavior with different $SiO₂/$ Al_2O_3 ratios suggests that differences in the nature or the extent of the defects in the proposed structure exist for allophanes of different chemical compositions. Tables 2 and 3 show that both the intensity of the exothermic peak and the temperature at which mullite was was detected were higher for H-clays with high $SiO₂/$ Al_2O_3 ratios than those with low ratios. This may be explained by a smaller defect structure of kaolinite, namely better crystallinity or longer range order in the structure of allophanes with high ratios than those with low ratios. The strength of adsorbed water or structural OH supports the hypothesis that allophane with high $SiO₂/Al₂O₃$ ratios are less poorly crystalline than allophane with low $SiO₂/Al₂O₃$ ratios.

Properties, such as surface acidity, that are intimately connected with the structure were reported to be dependent on the SiO_2/Al_2O_3 ratio (Henmi, 1977). Different structure models have been proposed for allophanes having $SiO₂/Al₂O₃$ ratios of 1 and 2 (Okada *et al.,* 1975; Wada, 1978). The differences in such properties and structures among allophanes of different chemical compositions might be discernible by means of ordinary mineralogical analyses, such as XRD and IR, although these differences would not be large in comparison with the differences found between allophane and other related clay minerals, such as imogolite and halloysite. These differences might be useful in dividing allophane into subspecies.

ACKNOWLEDGMENTS

The author is greatly indebted to Dr. R. L. Parfitt, Soil Bureau, New Zealand Department of Scientific and Industrial Research, for his kind and critical reading of this manuscript.

REFERENCES

- Barshad, I. (1965) Thermal analysis techniques for mineral identification and mineralogical composition: in *Methods of Soil Analysis, Part 1 Physical and Mineralogical Properties, Including Statistics of Measurement and Sampling, C.* A. Black, ed., *Amer. Soc. Agron.,* 699-742.
- Brindley, G. W. and Fancher, D. (1969) Kaolinite defect structure; possible relation to allophane: *Proc. Int. Clay Conf. 1969* 2, 29-34.
- Campbell, A. S., Mitchell, B. D., and Bracewell, J. M. (1968) Effect of particle size, pH and organic matter on the thermal analysis of allophane: *Clay Miner.* 7, 451-454.
- Egawa, T. (1964) A study on coordination number of aluminum in allophane: *Clay Sci.* 2, 1-7.
- Fieldes, M. (1955) Clay mineralogy of New Zealand soils, Part II: allophane and related mineral colloids: *N. Z. J. Sci. Technol.* B 37, 336-350.
- Fieldes, M. (1957) Clay mineralogy of New Zealand soils, Part 4: differential thermal analysis: *N. Z. J. Sci. Technol.* B 38, 533-570.
- Fieldes, M. and Claridge, G. G. C. (1975) Allophane: in *Soil Components Vol. 2 Inorganic Components,* J. E. Gieseking, ed., Springer-Verlag, New York, 351-393.

Henmi, T. (1977) The dependence of surface acidity on chem-

ical composition $(SiO_2/Al_2O_3$ molar ratio) of allophanes: *Clay Miner.* 12, 356-358.

- Henmi, T. and Wada, K. (1976) Morphology and composition of allophane: *Amer. Miner.* 61, 379-390.
- Higashi, T. and Ikeda, H. (1974) Dissolution of allophane by acid oxalate solution: *Clay Sci.* 4, 205-212.
- Jackson, M. L. (1956a) Dispersion of soil minerals: in *Soil Chemical Analysis--Advanced Course,* Published by the author, 31-95.
- Jackson, M. L. (1956b) Differential thermal analysis of soil minerals: in *Soil Chemical Analysis--Advance Course*, Published by the author, 251-329.
- Kitagawa, Y. (1971) The "unit particle" ofallophane: *Amer. Miner. 56,* 465-475.
- Mackenzie, R. C. (1970) Simple phyllosilicates based on gibbsite- and brucite-like sheets: in *Differential Thermal Analysis Vol.* 1, R. C. Mackenzie, ed., Academic Press, London, 498-537.
- Mackenzie, R. C. and Mitchell, B, D. (1970) Technique: **in** *Differential ThermalAnalysis Vol.* 1, R. C. Mackenzie, ed., Academic Press, London, 101-122.
- Mehra, O. P. and Jackson, M. L. (1960) Iron oxide removal from soils and clays by a dithionite-citrate system with sodium bicarbonate buffer: *Clays & Clay Minerals* 7, 317-327.
- Mitchell, B. D., Farmer, V. C., and McHardy, W. J. (1964) Amorphous inorganic materials in soils: *Adv. Agron.* 16, 327-383.
- Miyauchi, N. and Aomine, S. (1966) Effect of exchangeable cations on the high temperature exothermic peak of allophane: *Soil Sci. Plant Nutr.* 12, 13-17.
- Miyazawa, K. (1966) Clay mineral composition of andosoils in Japan--with reference to their classification: *Bull. Nat. Inst. Agr. Sci.* B 17, 1-100.
- Okada, K., Morikawa, S., Iwai, S., Obira, Y., and Ossaka, J. (1975) A structural model ofallophane: *Clay Sci.* 4,219- 303.
- Ossaka, J. (1960) On the hydro-alumina silicate minerals from Mt. Asama: *Adv. Clay Sci.* 2, 339-349.
- Ossaka, J. (1961) On the mechanism of formation of precipitated allophane: *Adv. Clay Sci.* 3, 225-233.
- Ossaka, J. (1962) On the heat transformation of hydrated low crystalline materials from system SiO₂-Al₂O₂: *Adv. Clay Sci.* 4, 33-47.
- Shoji, S. and Saigusa, M. (1977) Amorphous clay materials of Towada Andosoils: Soil Sci. Plant Nutr. 23, 437-455.
- Udagawa, S., Nakada, T., and Nakahira, M. (1969) Molecular structure of allophane as revealed by its thermal transformation: *Proc. Int. Clay Conf. (Tokyo) Vol. 1,* 151-159.
- Wada, K. (1977) Allophane and imogolite: in *Minerals in Soil Environments,* J. B. Dixon and S. B. Weed, eds., Soil Sci. Soc. Amer., 603-638.
- Wada, K. (1978) Structural formulas of allophanes: *Proc. Int. Clay Conf. (Oxford),* 537-545.
- Wada, K. and Harward, M. E. (1974) Amorphous clay constituents of soils: *Adv. Agron.* 26, 211-260.
- Wada, K., Henmi, T., Yoshinaga, N., and Patterson, S. H. (1972) Imogolite and allophane formed in saprolite of basalt on Maul, Hawaii: *Clays & Clay Minerals* 20, 375-380.
- Wada, K. and Wada, S. (1976) Clay mineralogy of the B horizon of two Hydrandepts, a Torrox and a Humitropept in Hawaii: *Geoderma* 16, 139-I57.
- Wada, K. and Yoshinaga, N. (1969) The structure of imogolite: *Amer. Miner.* 54, 50-71.
- Wada, K., Yoshinaga, N., Yotsumoto, H., Ibe, K., and Aida, S. (1970) High resolution electron micrographs of imogolite: *Clay Miner.* 8, 487-489.
- Wada, S. and Wada, K. (1977) Density and structure of allophane: *Clay Miner.* 12, 289-298.
- Wells, N., Childs, C. W., and Downes, C. J. (1977) Silica

spring. Tongariro national park, New Zealand—analyses of the spring water and characterization of the alumino-silicate deposit: Geochem. Cosmochim. Acta 41, 1497-1506.

Yoshinaga, N. (1966) Chemical composition and some thermal data of eighteen allophanes from andosoils and weathered pumices: Soil Sci. Plant Nutr. 12, 47-54.

Yoshinaga, N., Yotsumoto, H., and Ibe, K. (1968) An electron microscopic study of soil allophane with an ordered structure: Amer. Miner. 53, 319-323.

(Received 18 July 1979; accepted 11 November 1979)

Резюме-Были обнаружены различия в кривых дифференциального термического анализа и в температурах развития новой фазы между аллофанами высоких (1,91-1,99) и низких (1,47-1,53) отношений SiO₂/Al₂O₃. Эндотермический пик из-за постоянной дегидротации и дегидроксилирования наблюдался при высоких температурах (153°–185°С) для аллофанов с высокими отношениями SiO_2/AI_2O_3 и при низких температурах (148°-165°С) для аллофанов с низкими отношениями SiO_2/A Al₂O₃. Для аллофанов с высокими отношениями температура эндотермического пика была ниже и высота более обусловлена содержанием обменных катионов, чем для аллофанов с низкими отношениями. Новые фазы не развились в аллофанах с высокими отношениями SiO_9/Al_2O_3 даже после подогрева до 1000°С, выше температуры экзотермического пика. Наоборот, симптоматическое развитие новых фаз было замечено в аллофанах с низкими отношениями SiO₂/Al₂O₃ при 900°С, ниже температуры экзотермического пика. Влияние отношения SiO_2/AI_2O_3 на термическое поведение аллофанов указывает на то, что различия в структуре тесно связаны с химическим составом этого материала. [N. R.]

Resümee-Bei den DTA-Kurven und bei den Bildungstemperaturen neuer Phasen wurden Unterschiede zwischen Allophanen mit hohen (1,91 bis 1,99) und niedrigen (1,47 bis 1,53) SiO₂/Al₂O₃-Verhältnissen gefunden. Der endotherme Peak, verursacht durch die ständige Dehydrierung und Dehydroxylierung, war für Allophane mit hohem SiO_2/Al_2O_3 -Verhältnis bei höheren Temperaturen (153° bis 195°C). Für Allophane mit niedrigem SiO₂/Al₂O₃-Verhältnis war er bei niedrigeren Temperaturen (148° bis 165°C). Bei Allophanen mit großen Verhältniszahlen war die Temperatur des endothermen Peaks niedriger, und die Höhe mehr durch die austauschbaren Kationen beeinflußt, als bei solchen mit kleinen Verhältniszahlen. In Allophanen mit hohen SiO₂/Al₂O₃-Verhältnissen wurden keine neuen Phasen gebildet, selbst wenn sie auf über 1000°C, über die Temperatur des exothermen Peaks, erhitzt wurden. Im Gegensatz dazu wurde bei Allophanen mit niedrigen SiO₂/Al₂O₃-Verhältnissen die symptomatische Bildung neuer Phasen bei 900°C, unter der Temperatur des exothermen Peaks, beobachtet. Die Wirkung des SiO_2/AI_2O_3 -Verhältnisses auf das thermische Verhalten von Allophan legt die Annahme nahe, daß Unterschiede in der Struktur eng mit der chemischen Zusammensetzung dieses Materials zusammenhängen. [U. W.]

Résumé—Des différences ont été trouvées dans les courbes d'analyse thermale différentielle et dans les températures de développement de nouvelle phase entre les allophanes aux proportion hautes (1,91-1,99) et basses (1,47-1,53) de SiO_2/AI_2O_3 . Le sommet endothermique dû à la déshydration et à la déshydroxylation continuelles était à des températures plus hautes (153°–185°C) pour des allophanes avec des proportions élevées de SiO₂/Al₂O₃ et à de plus basses températures (148[°]-165[°]C) pour celles avec les proportions les plus basses. De nouvelles phases ne sont pas développées dans les allophanes ayant des proportions élevées de SiO₂/Al₂O₃, même après échauffement à 1000°C, au-delà de la température du sommet exothermique. Par contraste, un développement symptomatique de nouvelles phases a été noté dans les allophanes ayant de basses proportions de SiO_2/Al_2O_3 à 900°C, en dessous de la température du sommet exothermique. L'effet de la proportion de $SiO₂/Al₂O₃$ dans le comportement thermal de l'allophane suggère fortement que les différences dans la structure sont associées de près avec la composition chimique de ce matériel. [D. J.]