

## Mineral Analyses of Extraterrestrial Metal

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Metal is observed in almost all types of meteorites (irons, stony-irons, and stony) and in lunar soils. The metal contains Fe with varying amounts of Ni (~5 to 50 wt%) and minor amounts of Co, P, S, and C. Groups of iron meteorites, classified by their trace element contents, are thought to crystallize in the cores of asteroids (protoplanets) while other groups of meteorites crystallize in the silicate in the outer portion of such bodies [1]. The cooling rates of the iron meteorites in the temperature range 500 to 600°C vary greatly from >10,000 to <1°C/million years (°C/My). These slow cooling rates allow for phase transformations to take place which cannot be duplicated in the laboratory and should approach equilibrium. The objective of this paper is to show how advanced instrumentation is used to characterize the minerals that form from these transformations.

Fig. 1 shows the Fe-Ni phase diagram. A typical iron meteorite of 10wt% Ni will cool as a single crystal from fcc taenite into the bcc kamacite ( $\alpha$ ) + taenite ( $\gamma$ ) phase field around 700°C forming a Widmanstätten pattern (Fig. 2). As cooling proceeds, a Ni gradient develops in the taenite as Ni moves into the host phase from the growing kamacite. At lower temperatures, the remaining taenite transforms through a martensite reaction to a two phase assemblage of  $\alpha$  +  $\gamma$  called plessite and high Ni taenite transforms by a spinoidal reaction to a two phase assemblage of fcc  $\gamma''$  and  $\gamma_1$ (cloudy zone). Fig. 3a shows a fcc EBSD map of the Carlton meteorite in which the high Ni rim of the  $\gamma$  and the taenite in the plessite have the same orientation as the parent taenite single crystal. Orientation maps (Fig 3b) confirm that the orientation of the bcc phase to the fcc phase is Kurdjumov Sachs [2], (110) bcc // (111) fcc.

Thin foils of IVA irons were prepared for TEM using dual beam FEI DB-235 focused ion beam (FIB) instruments. Selected  $\alpha/\gamma$  interface regions were analyzed using a FEI Tecnai F30ST field emission transmission - analytical electron microscope (TEM-AEM) at Sandia National Laboratories. Fig. 4 shows the Ni gradient in the taenite phase close to the kamacite-taenite boundary in Carlton obtained using high resolution x-ray microanalysis. The high Ni region varies from about 55 wt% Ni to about 40 wt% Ni. A two phase structure is observed in the Ni x-ray map (Fig 4b) at lower Ni contents. Below 40 wt% Ni, a spinodal structure (fcc  $\gamma''$  and  $\gamma_1$ (cloudy zone) with the high Ni (~50wt%)  $\gamma''$  as the round phase and the low Ni (~5 wt%)  $\gamma_1$  as the matrix phase) has formed. Measurements of the Ni gradient in taenite, microns beyond the  $\alpha$ - $\gamma$  interface, allow for the measurement of the cooling rate of the Carlton meteorite, between 8 and 45 °C/My [3]. The Carlton iron also contains significant C >0.2wt%, and (Fe-Ni)<sub>3</sub>C and (Fe-Ni)<sub>23</sub>C<sub>6</sub> form on cooling. C is also distributed between kamacite and taenite. We used a Cameca IMS 1280 ion probe with a Cs<sup>+</sup> primary beam to measure the C content. The spatial resolution is 5-7 $\mu$ m and the C detection limit is about 1ppm, .0001 wt%. Carlton contains 12  $\pm$  4 ppm C in kamacite and 330 to 400 ppm C in the high Ni taenite region. Equilibration from high temperature to at least 400°C is confirmed. Mineral analysis at 1-2nm spatial resolution using the TEM-AEM remains challenging due to the difficulty of making very thin specimens, <25nm.

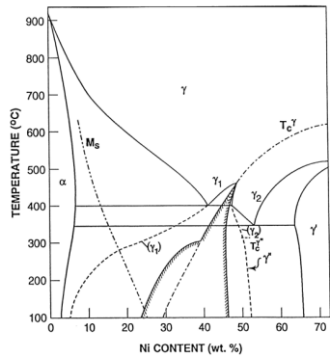


Fig 1. Fe-Ni Phase Diagram Kamacite – α, Taenite – γ Dashed line is the spinodal

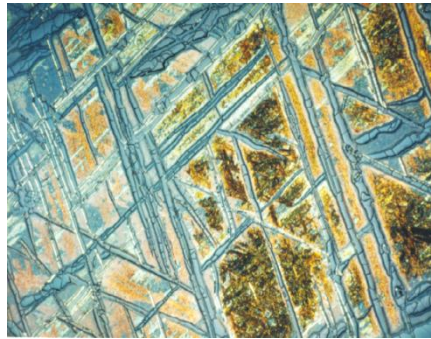


Fig. 2 Carlton iron meteorite Widmanstätten pattern [1]. Kamacite plates are blue.

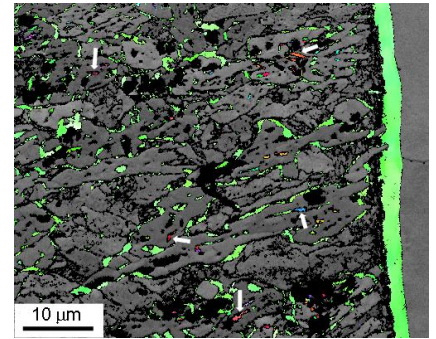


Fig. 3a EBSD fcc orientation map of Carlton. Green regions have the same orientation as the γ parent single crystal [2]

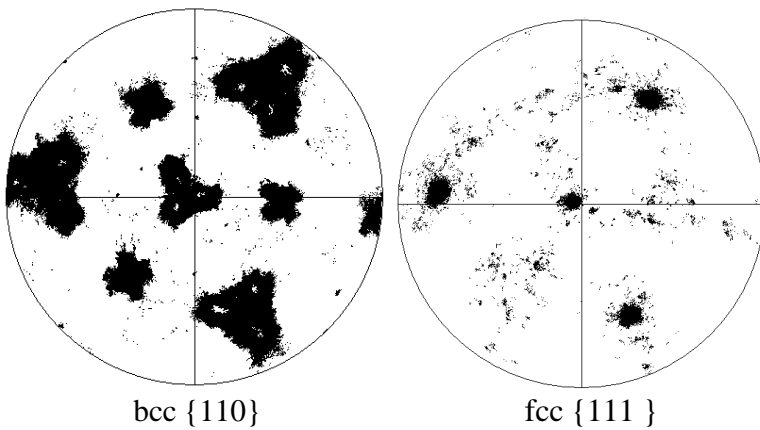


Fig. 3b. Pole figures obtained from EBSD map, Fig 3a

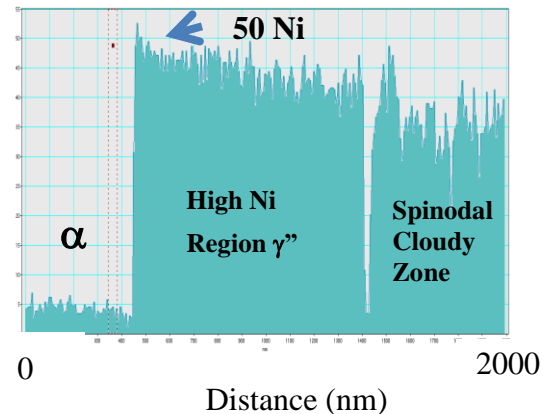


Fig. 4a Ni concentration vs. distance of the outer taenite rim of the Carlton meteorite

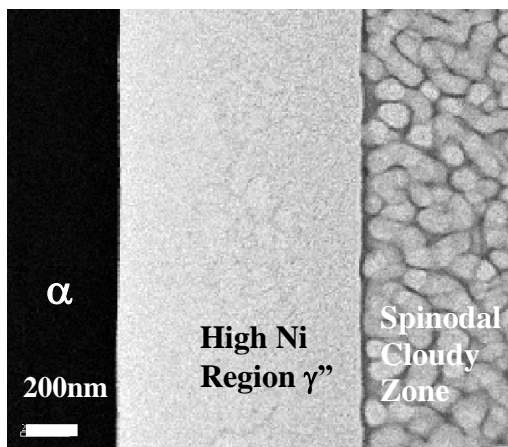


Fig 4b. 2x2 μm Ni x-ray map of the outer taenite rim of the Carlton meteorite.

References  
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 [3] JI Goldstein et al., *Geochem. Cosmochem. Acta.*, submitted (2014)  
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